

Geologic and geomorphic impacts of the 2010-2012 Canterbury earthquake sequence and local evidence for large prehistoric earthquakes

M.C. Quigley

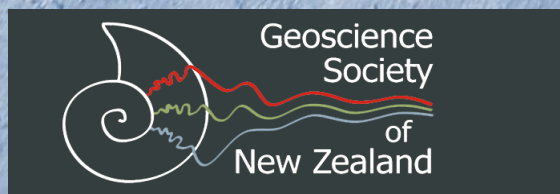
Department of Geological Sciences

University of Canterbury

2013 Hochstetter Lecture

THANKS!

Funders



Collaborators



Students

Dr. Brendan Duffy, Timothy Stahl, Dr. Eric Bilderback, Narges Khajavi, Sarah Bastin, Sharon Hornblow, Gregory De Pascale, Peri Sasnett, and many others!

A decade of blissful seismic quiescence beneath the Canterbury Plains

1995 M 6.0

1995 Cass M 6.2

1994 AP M 6.7

1994 M 6.0

1946 M 6.2

No obvious 'precursory' seismicity at future site of CES

A history of earthquake clustering
Highest decadal seismicity rates in the region near location of largest recent earthquakes (aftershocks?)

1869 M_w 4.8

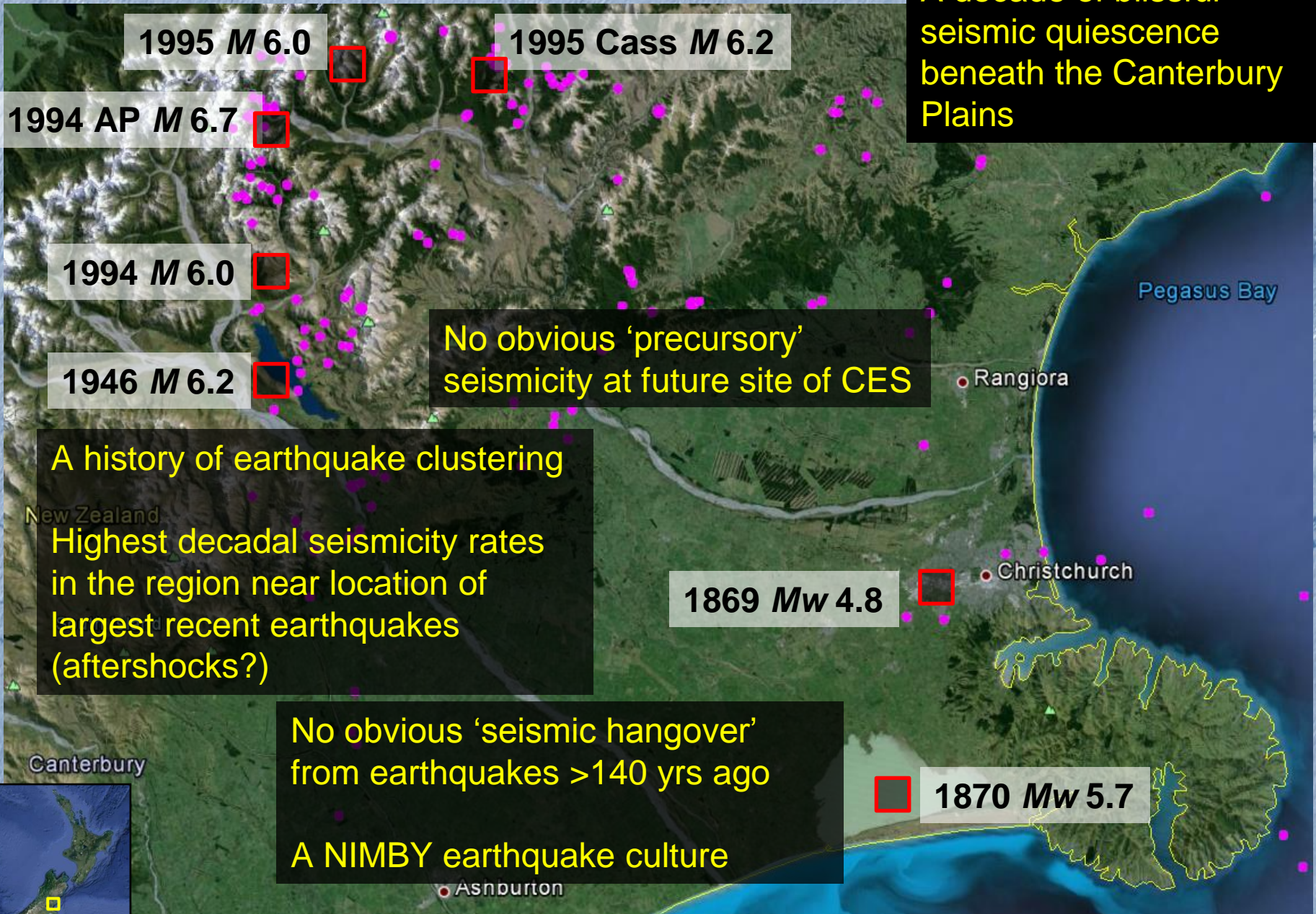
No obvious 'seismic hangover' from earthquakes >140 yrs ago

A NIMBY earthquake culture

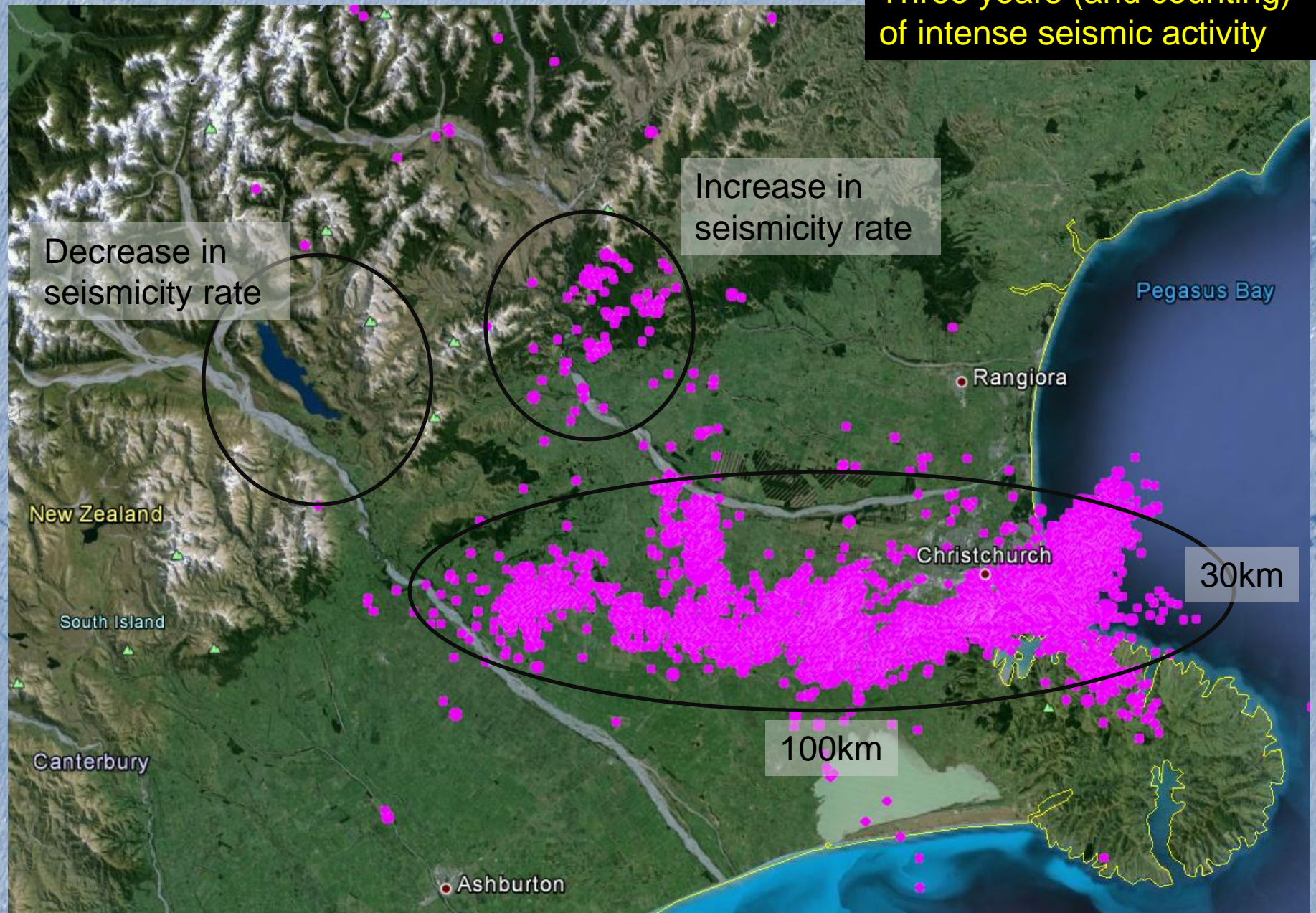
1870 M_w 5.7

Data source: Geonet

Seismicity Sept 1, 2000 to Sept 3, 2010
 $M \geq 3.0$, 0-15 km depth

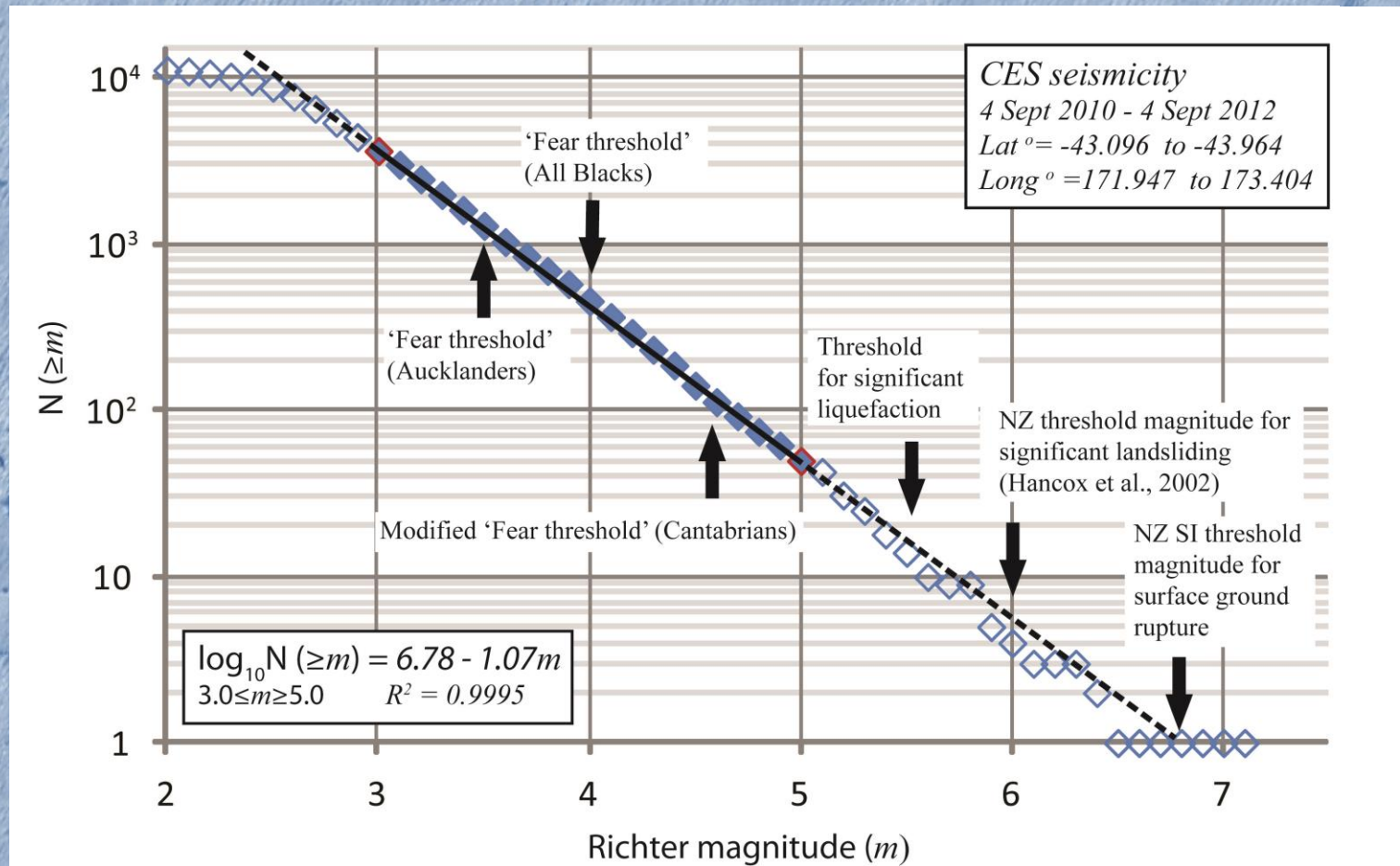


Three years (and counting)
of intense seismic activity



Seismicity Sept 4, 2010 to July 1, 2013
M \geq 3.0, 0-15 km depth

Truth and beauty of earthquakes: The Gutenberg-Richter relationship



1 earthquake sequence

100s of 'heart-in-throat' moments (prolonged stress and anxiety)

10 liquefaction episodes at some locations (4 major)

5 rockfall episodes at some sites (3 major)

Earthquake comparisons: Counting the costs

	4 September 2010	22 February 2011	13 June 13 2011	23 December 2011
Mag (M_w)	7.1	6.2	6.0	5.9
Epicentre¹	30 km W	10 km SE	10 km SE	10 km E
Time²	4:36 am	12.51 pm	2.20 pm	3.18 pm
Max PGA³	0.6g (0.3g CBD)	2.2g (0.8g CBD)	2.2g (0.4g CBD)	0.96g ⁴ (0.25g CBD)
Casualties	0 fatalities	185 fatalities	0 fatalities	0 fatalities
Building Damage	To older brick & URM	All pre-1970s & several modern buildings with eccentric design	Further residential damage in Port Hills & already damaged CBD buildings	Minor, but several instances of progressive failure buildings
Liquefaction	Widespread in eastern suburbs	Extreme damage in many eastern Christchurch suburbs	Further damage in eastern Christchurch suburbs	Minor damage in eastern Christchurch suburbs
Cost⁵	4-5 billion	15-20 billion	c. 1.5 billion	c. 26 million

Loss of life and most damage occurred in an 'aftershock' on a previously unknown 'blind' fault
 Most fatalities in two building collapses – building stock performed well from life safety perspective but poorly from a 'post-event functionality' perspective

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Notes:

1. Epicentral distances are with respect to Christchurch CBD

More recent cost estimates exceed \$40 Billion – this is almost 30% of New Zealand's real GDP

Earthquake comparisons: Counting the costs

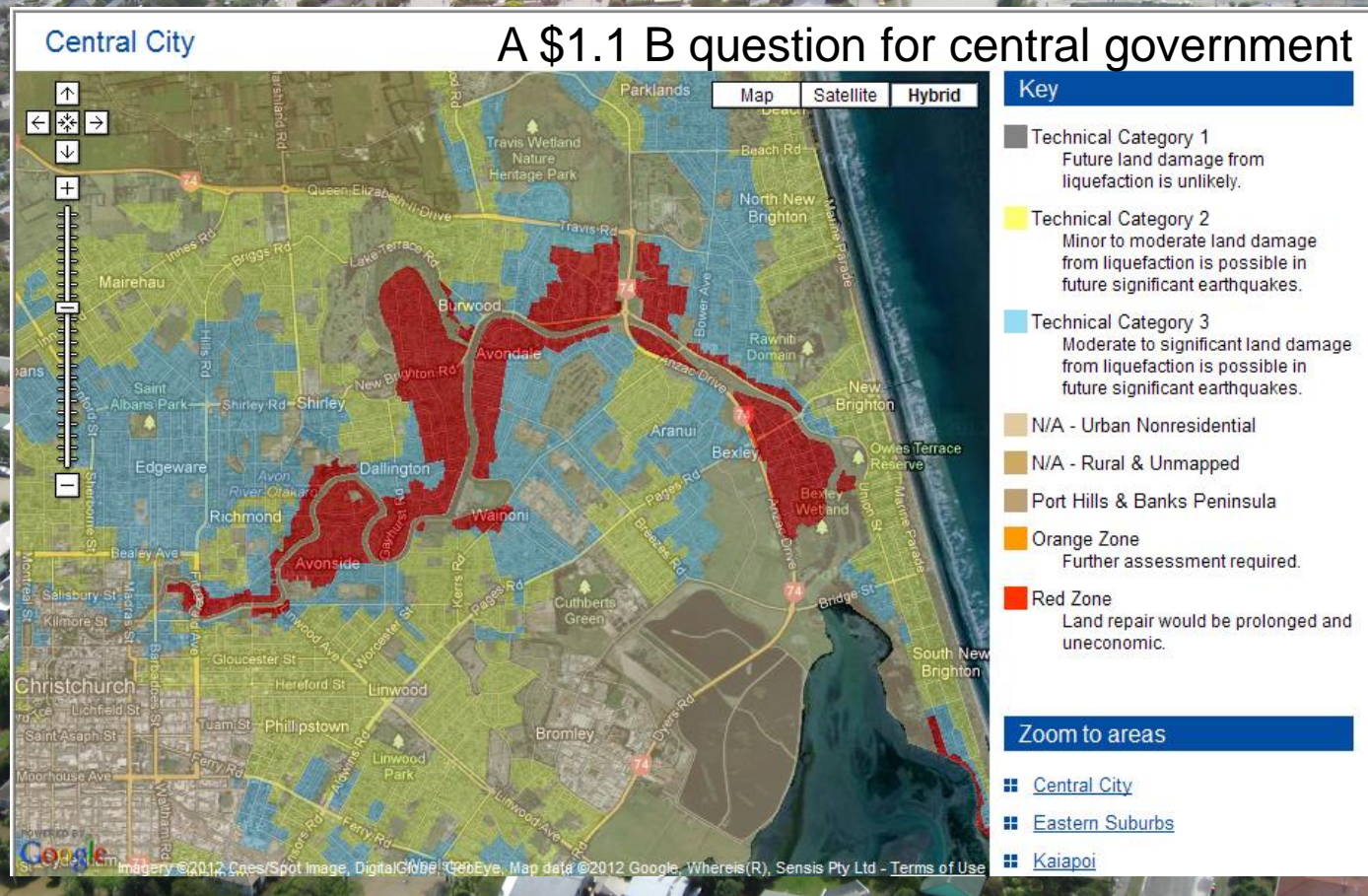
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Prob	1/475 yr	1/12,000 yr	1/1,000 yr	1/300 yr
%	0.2	0.008	0.08	0.3
Cost ⁵	4-5 billion	15-20 billion	c. 1.5 billion	c. 26 million

Notes:

1. Epicentral distances are with respect to Christchurch CBD

High PGAs and earthquake clustering
 Communicating science during a time-evolving hazard: The importance of discussing 'relative probability change' in addition to absolute probabilities

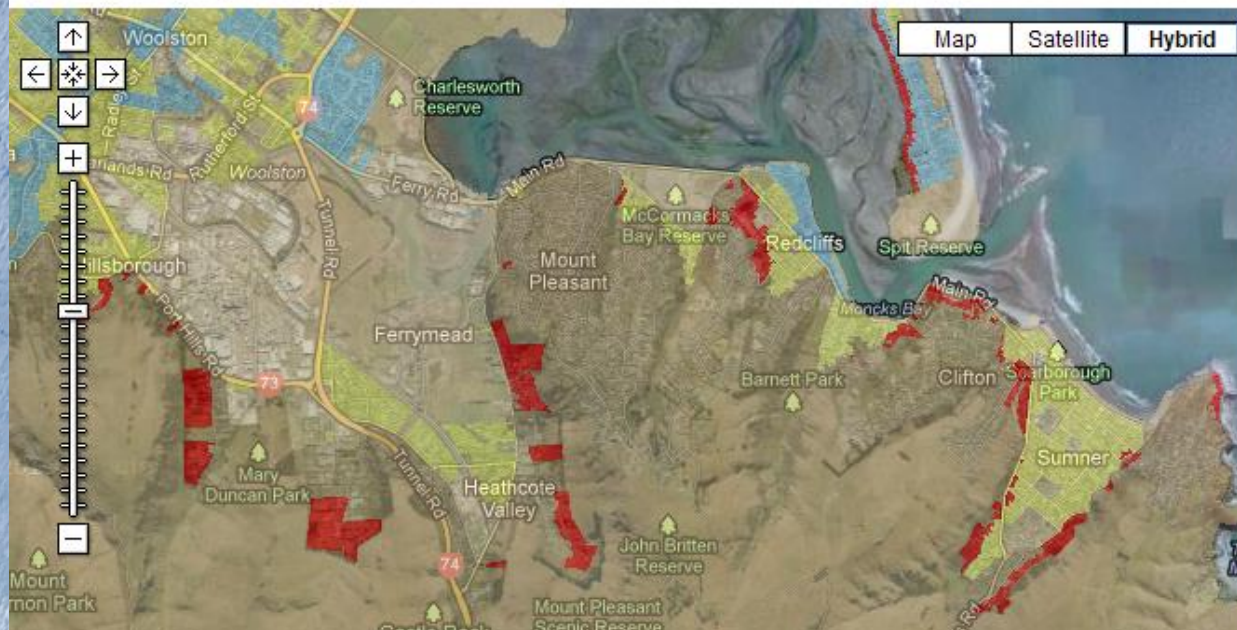
An acceptable risk or
an avoidable mistake?



Bexley: A modern suburb built at sea-level in a designated high-risk flood zone on ChCh's most liquefaction susceptible soils (1/75 to 1/100 yr threshold)



Central City



Key

- Technical Category 1
Future land damage from liquefaction is unlikely.
- Technical Category 2
Minor to moderate land damage from liquefaction is possible in future significant earthquakes.
- Technical Category 3
Moderate to significant land damage from liquefaction is possible in future significant earthquakes.
- N/A - Urban Nonresidential
- N/A - Rural & Unmapped
- Port Hills & Banks Peninsula
- Orange Zone
Further assessment required.
- Red Zone
Land repair would be prolonged and uneconomic.

Modern houses built immediately above and below seacliffs

5 rockfall fatalities, lots of luck, and some lessons learned

Some starting lessons

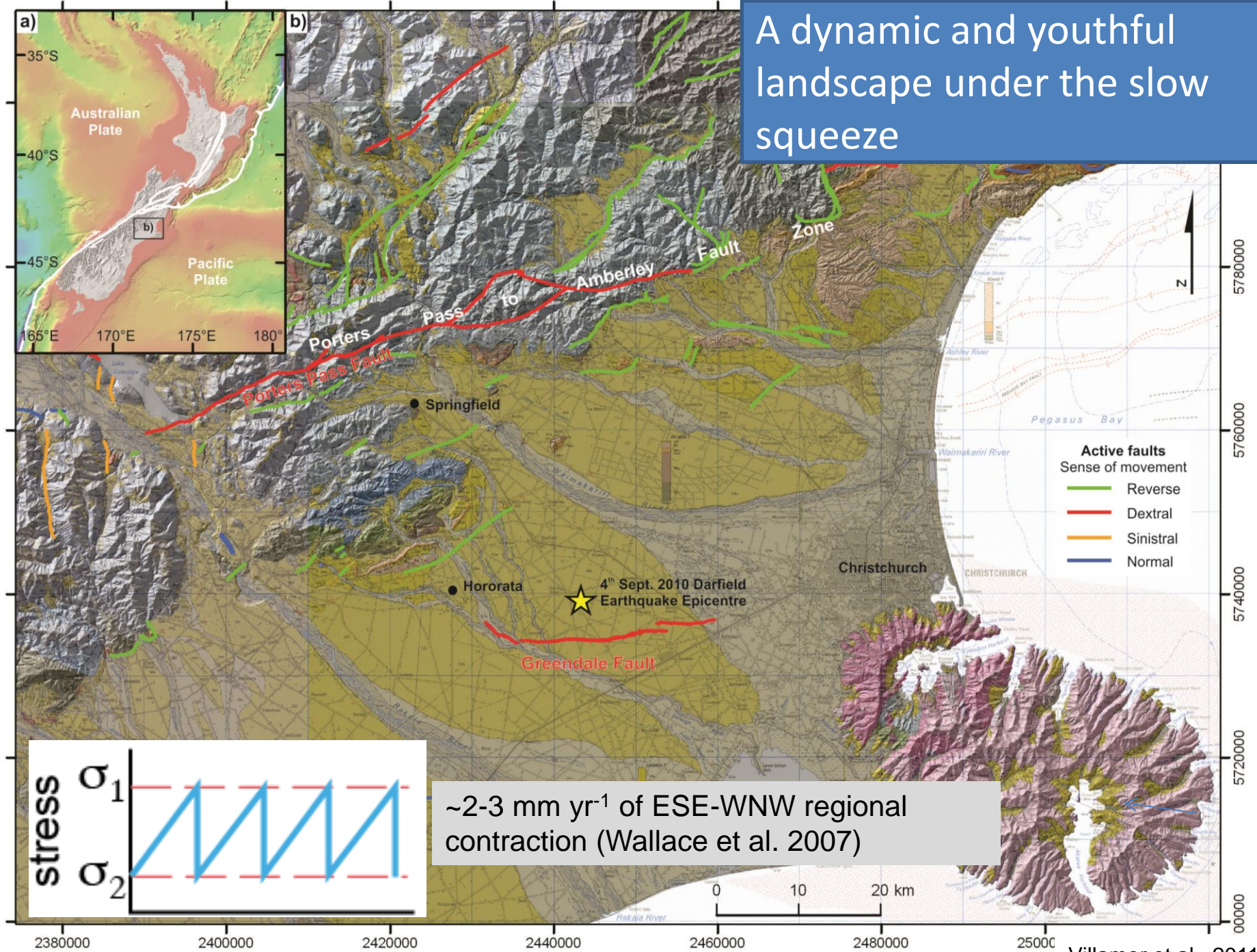
- Proactive, science-based land-use planning and structural and lifeline engineering is fundamental to reducing loss
- Reactive approach is more expensive, very complicated (science, politics, community well-being) and takes a large toll on people and the environment
- ‘Personalize your hazard’ (including ‘greenfields’)
- Combining ‘top-down’ and ‘bottom-up’ approaches to science communication will best facilitate good decision making and community acceptance

*Kaiapoi resident Brent Cairns says all he wants is transparency.
“I want to see is why my land deemed to be in the red zone, when we've
lived there for over a year.” TV3 Sept 2011*

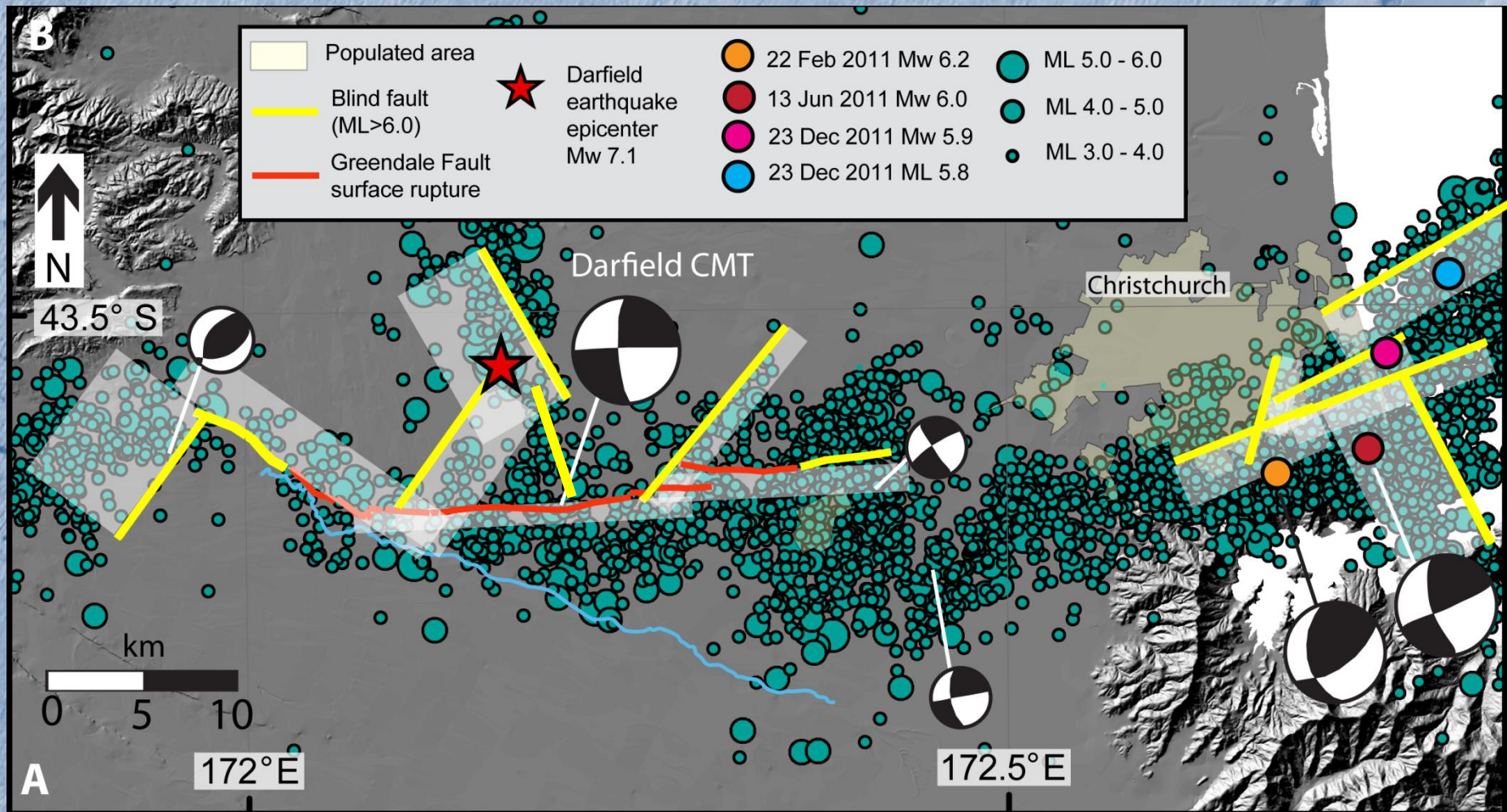
An aerial photograph of a dry, cracked landscape. The ground is covered in a network of deep, irregular cracks. A winding road or path runs through the lower right portion of the image. A small, dark, rectangular structure is visible near the center of the road. The overall color is a muted, dusty blue-grey.

Some science

A dynamic and youthful landscape under the slow squeeze

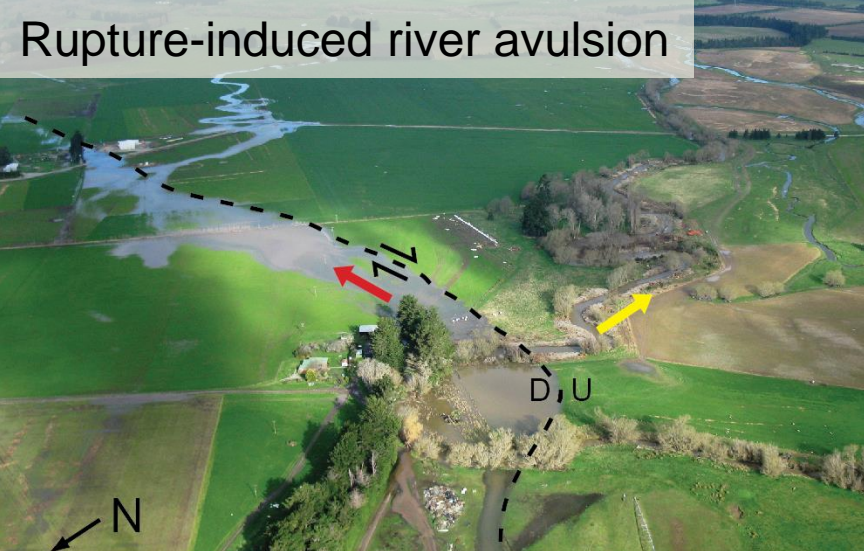


The 2010-2012 Canterbury earthquake sequence



Complex faulting (SS_D , SS_S , R, N)
1 surface rupture, at least 12 'blind' faults

Rupture-induced river avulsion



Damage to lifelines

Fault rupture damage: Important questions

- Relationship between earthquake magnitude, surface displacement, and SRL
- Thresholds between surface cracking and folding
- Width of deformation zone
- Return times (surface rupture and slip on related faults)

→ Forecasting earthquake hazards, designing resilient structures and lifelines, land-use planning (fault set-backs), interpreting paleo-earthquakes from the geologic record



Damage to structures

Surface rupture trace:
from the subtle to sublime



Photo: Mark Quigley



Photo: Russ Van Dissen

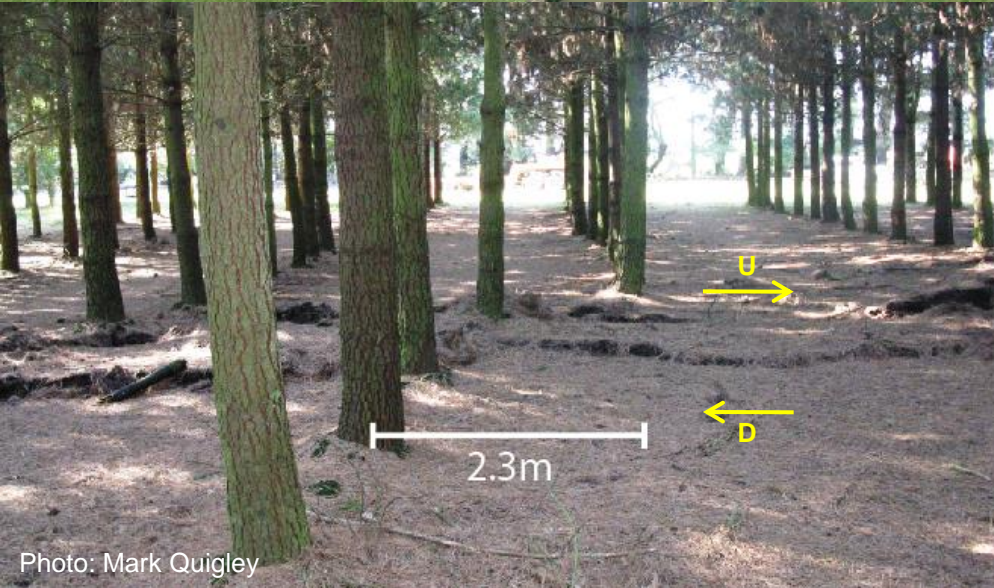


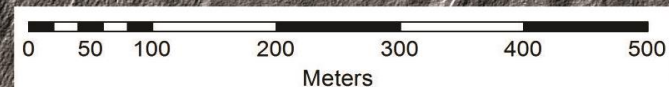
Photo: Mark Quigley



Photo: David Barrell

High resolution datasets

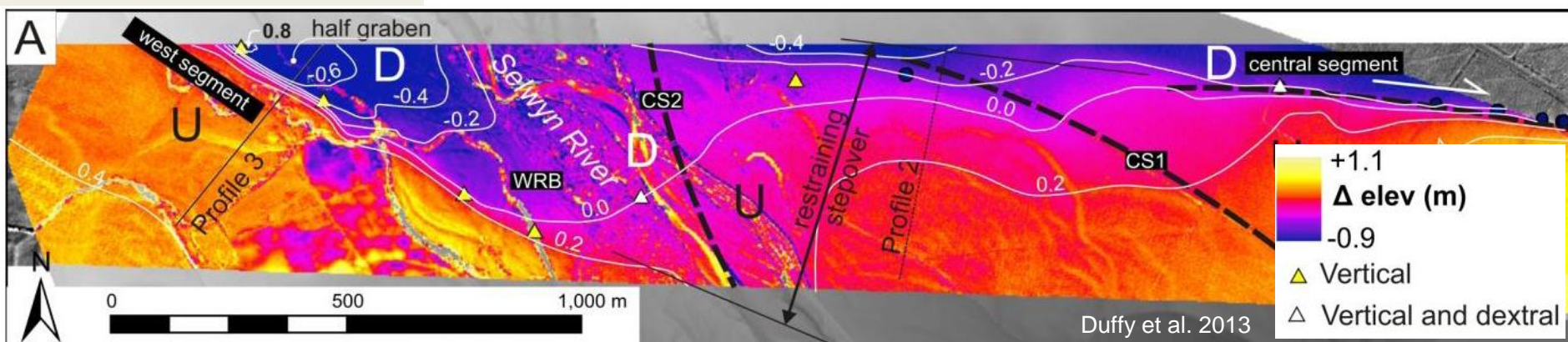
Airborne lidar



Terrestrial lidar

Lidar differencing

Courtesy Garth Archibald, GNS Science

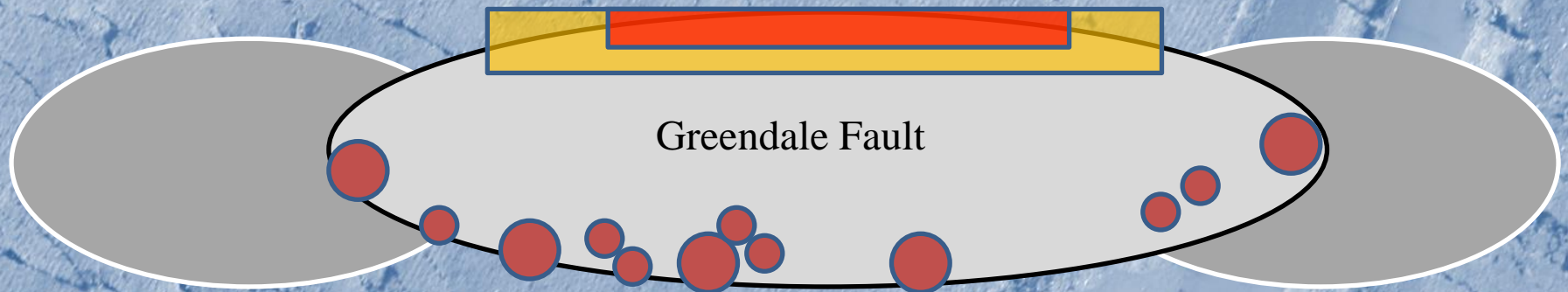


MAPPED SURFACE RUPTURE LENGTH (HISTORIC) = 29.5 ± 0.5 km

Mw 6.8-6.9

IDENTIFIABLE WITHOUT AGRICULTURAL FEATURES (GEOLOGIC) ≤ 20 km

Mw 6.6-6.7



GF SUBSURFACE RUP LENGTH (GEODETIC / SEISMOLOGIC) ~ 48 km

Mw 6.9-7.0

COMBINED SUBSURF RUP LENGTH (GEODETIC / SEISMOLOGIC) ~ 86 km

Mw 7.1

Importance of understanding how geologic record of active faulting rel to subsurface rupture potential:
 E Mw 7.1 = 6X E Mw 6.6

>100 field measurements,
1000s of potential strain markers



-Fault displacements
-Thresholds between
bending and breaking

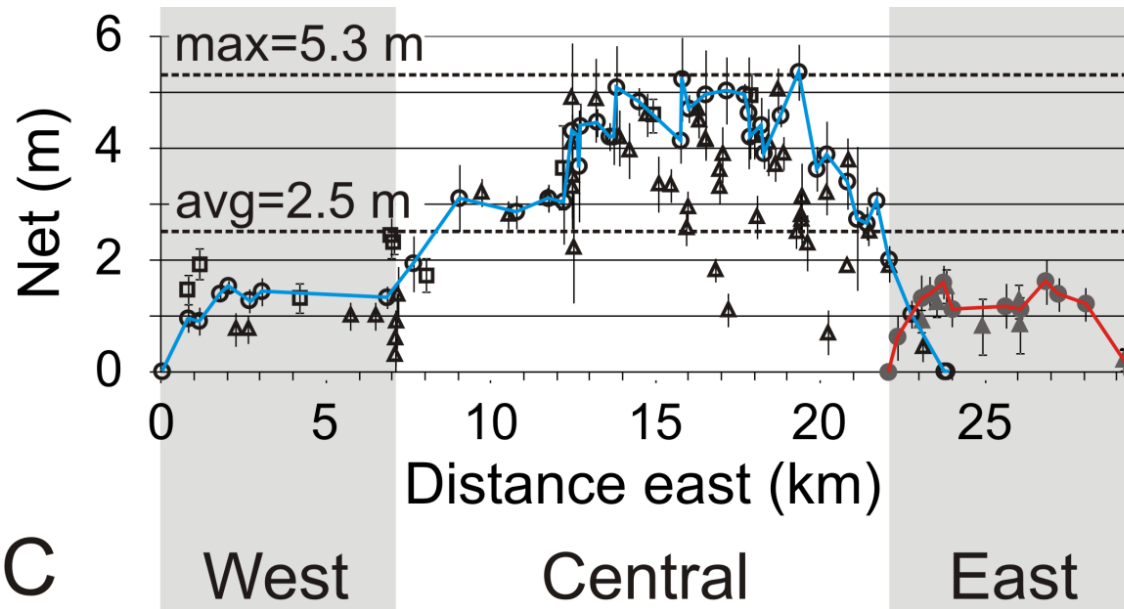
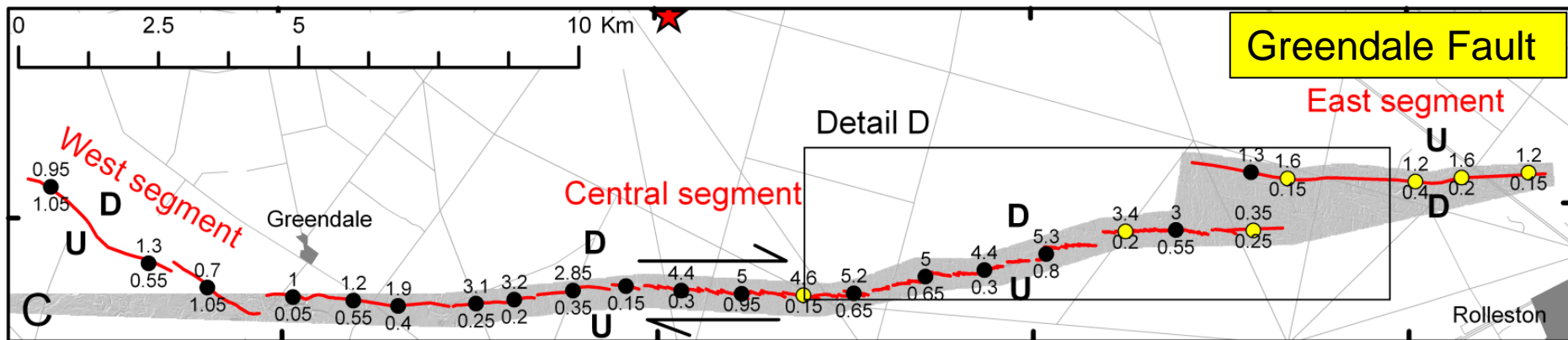


172°5'E

172°10'E

172°15'E

172°20'E



Mw 7.4 from
scaling rel

Why?

Lots of surface slip?
Slip distribution?
Measurement
technique?

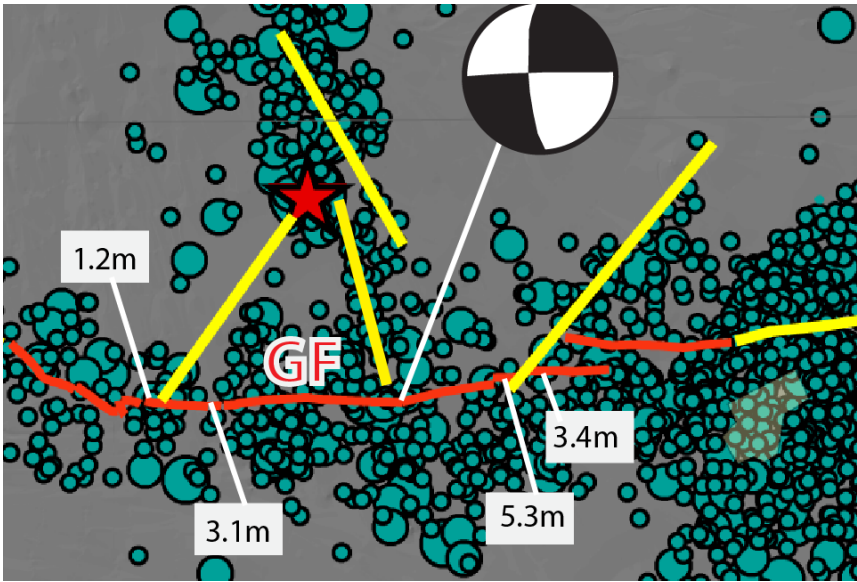
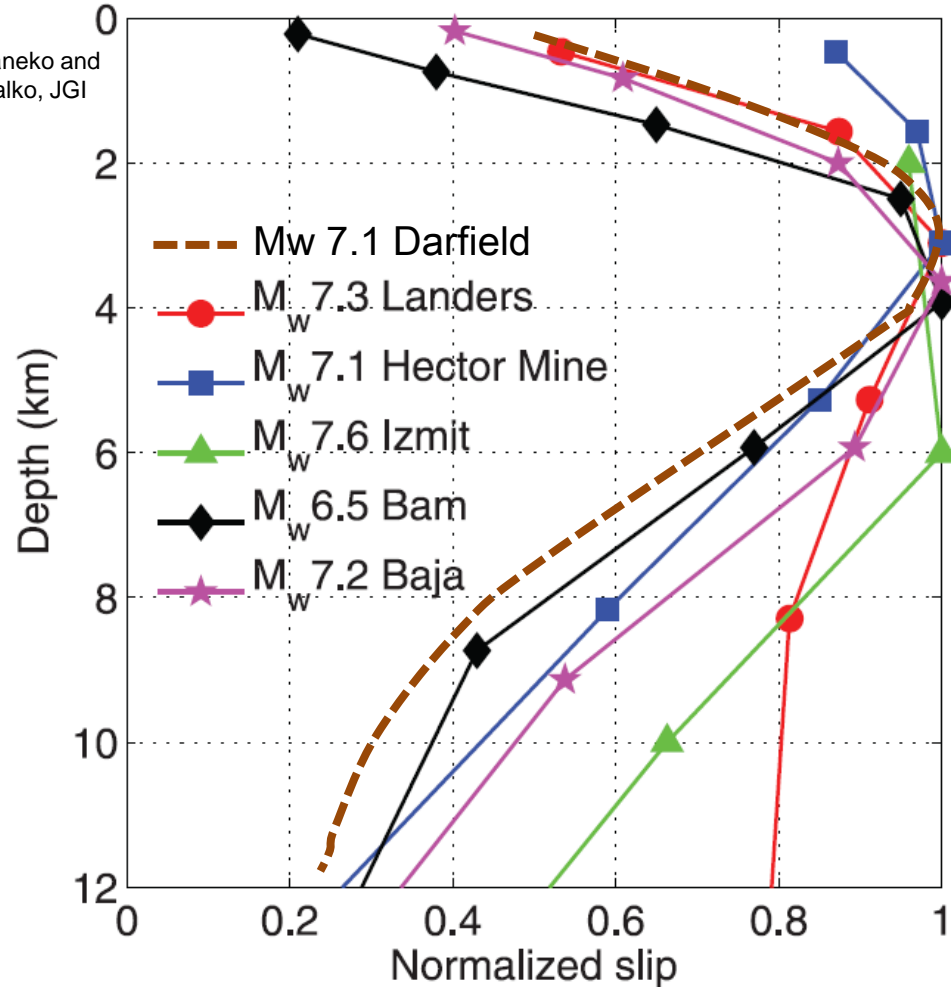
C

West

Central

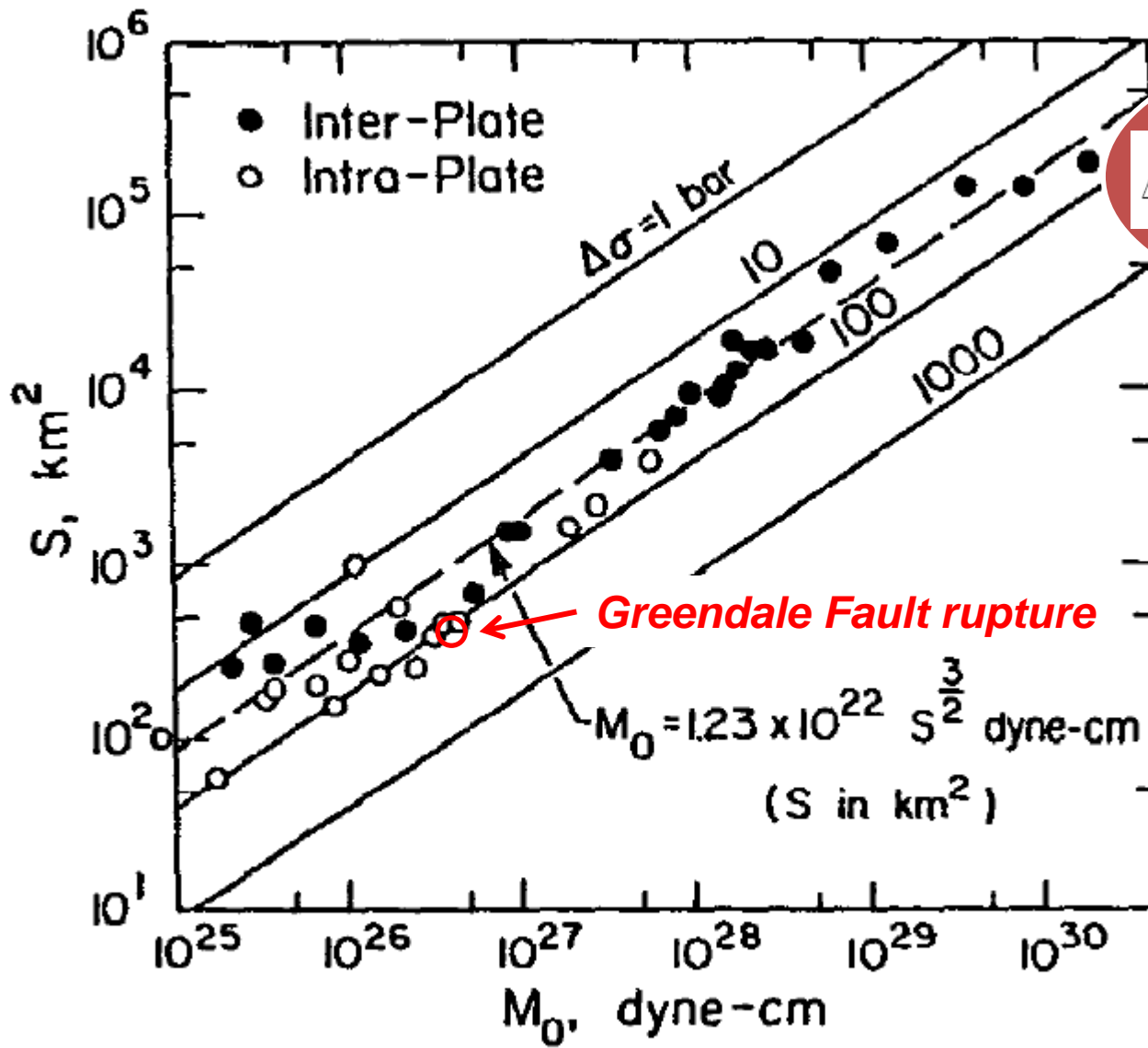
East

Quigley et al 2012



Some indication that
fault interactions may
have increased
coseismic slip

Nothing really anomalous
in slip distribution



$$\Delta\sigma^G = (\mu/C) \times (D/W) = 13.9 \pm 3.7 \text{ MPa}$$

At the high end of stress drop estimates but not surprising for tectonic setting

Higher than 'analogous' earthquakes ($8 \pm 1 \text{ MPa}$ for 1992 Mw 7.3 Landers; $10 \pm 2 \text{ MPa}$ for 1999 Mw 7.1 Hector Mine; Price and Bürgmann, 2002) = higher fault friction due to long recurrence intervals?

Relation between fault area and seismic moment for large and great earthquakes (Kanamori & Anderson 1975)

Quigley 2013



10m

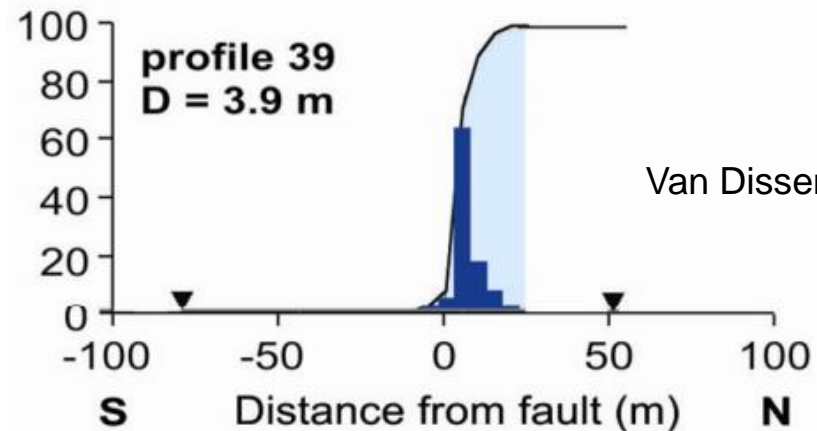
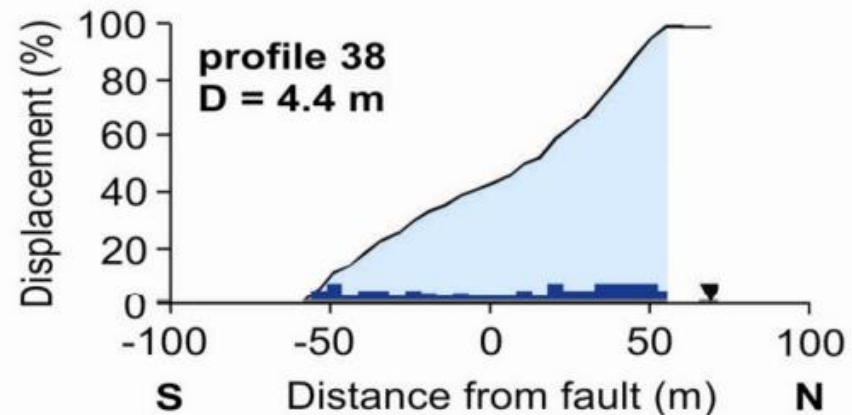
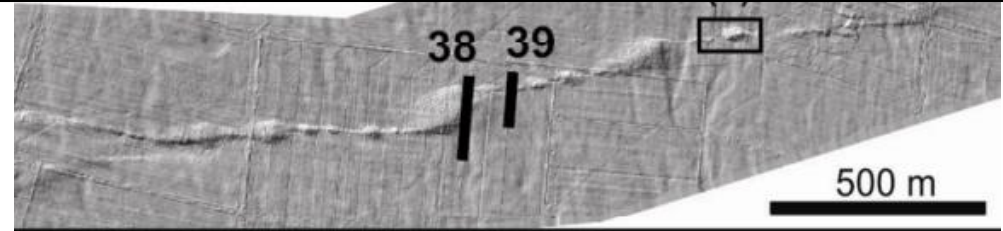
displacement on discrete fractures ~120cm

displacement across surface folding zone (fence not used) ~280cm

displacement from broad folding (using fences and agricultural features) ~340cm

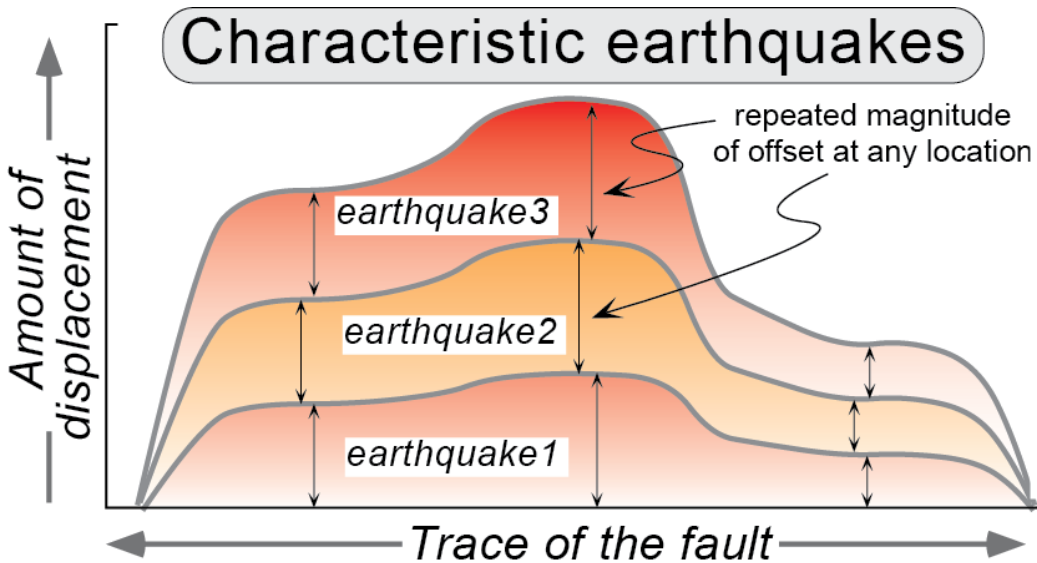
A

Better documentation of relationships between discrete and distributed deformation – Mw 7.0 estimated from discrete displacements only - confidence in eq scaling from geol offsets

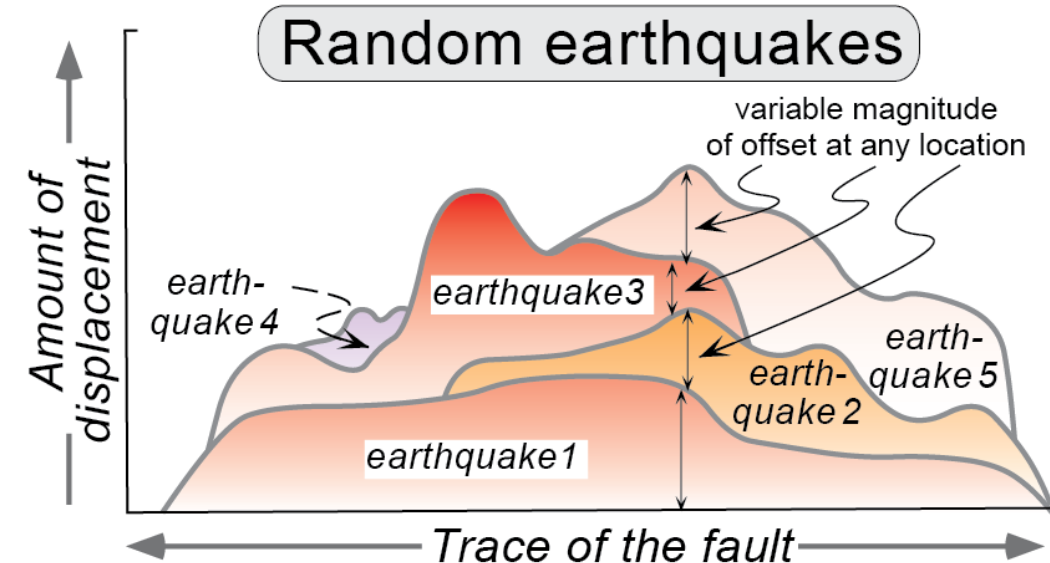


Van Dissen et al 2012

Greendale Fault behaviour in time and space

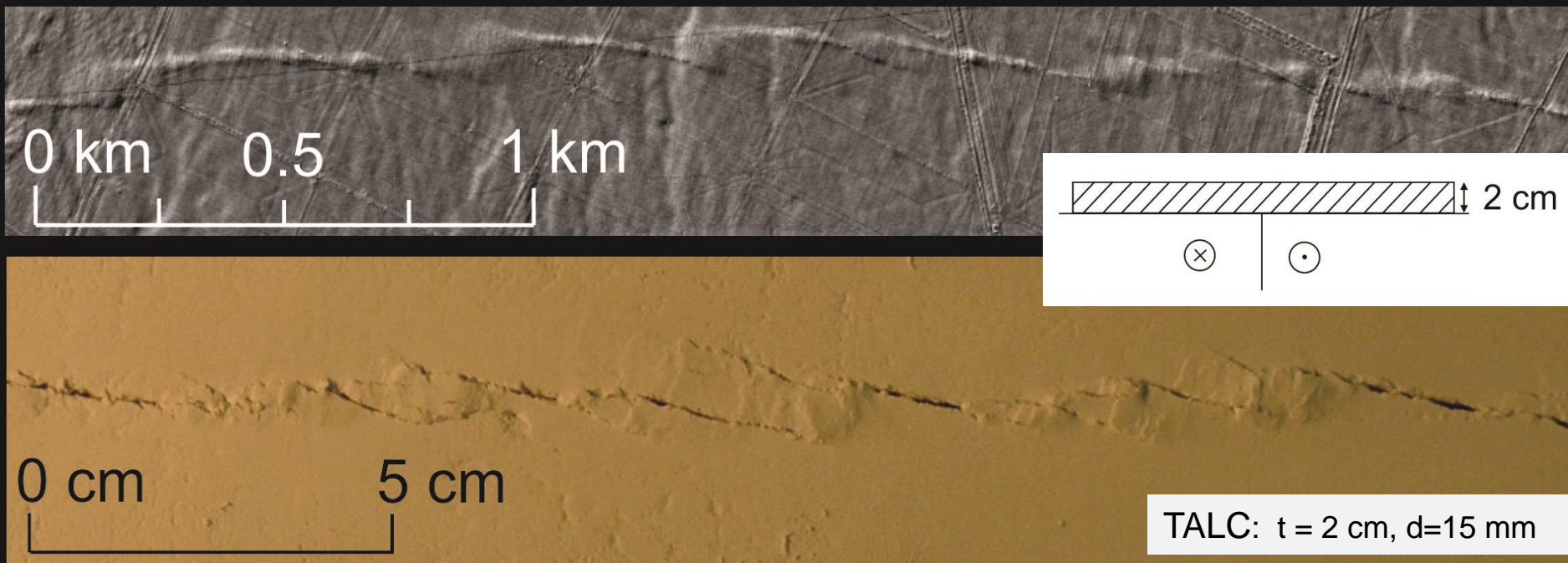


- Timing and M_w of penultimate eq
- RI, displacement histories, fault behaviour



Analogue modelling of Greendale Fault surface rupture:

What controls rupture morphology and displacement variations?
Where is the best place to site a trench, and what fractures will most faithfully record prior earthquakes?

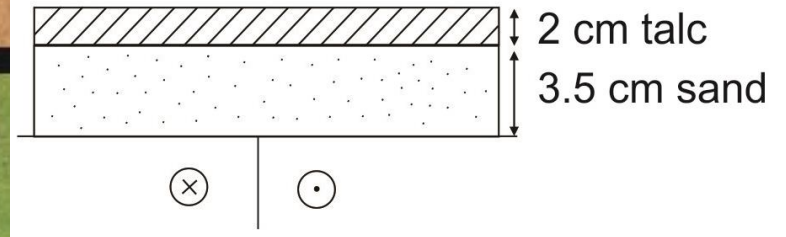


Single layer, cohesive material (talc) best replicates km-scale surface rupture morphology
Surface complexities created with simple, planar uniformly dipping basement fault

TALC_SAND: $d=19\text{ mm}$

0 cm 5 cm

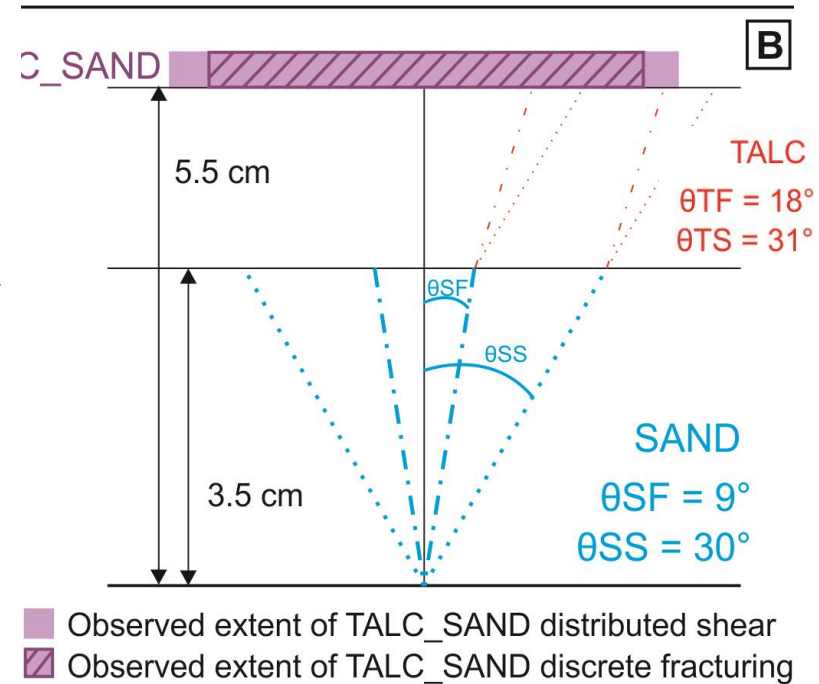
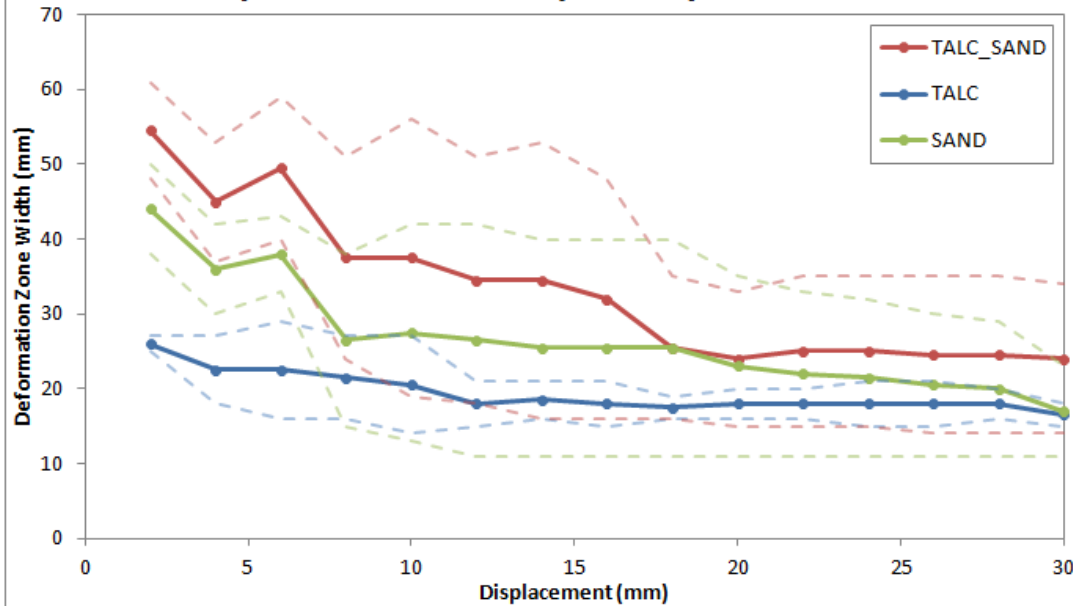
0 m 20 m



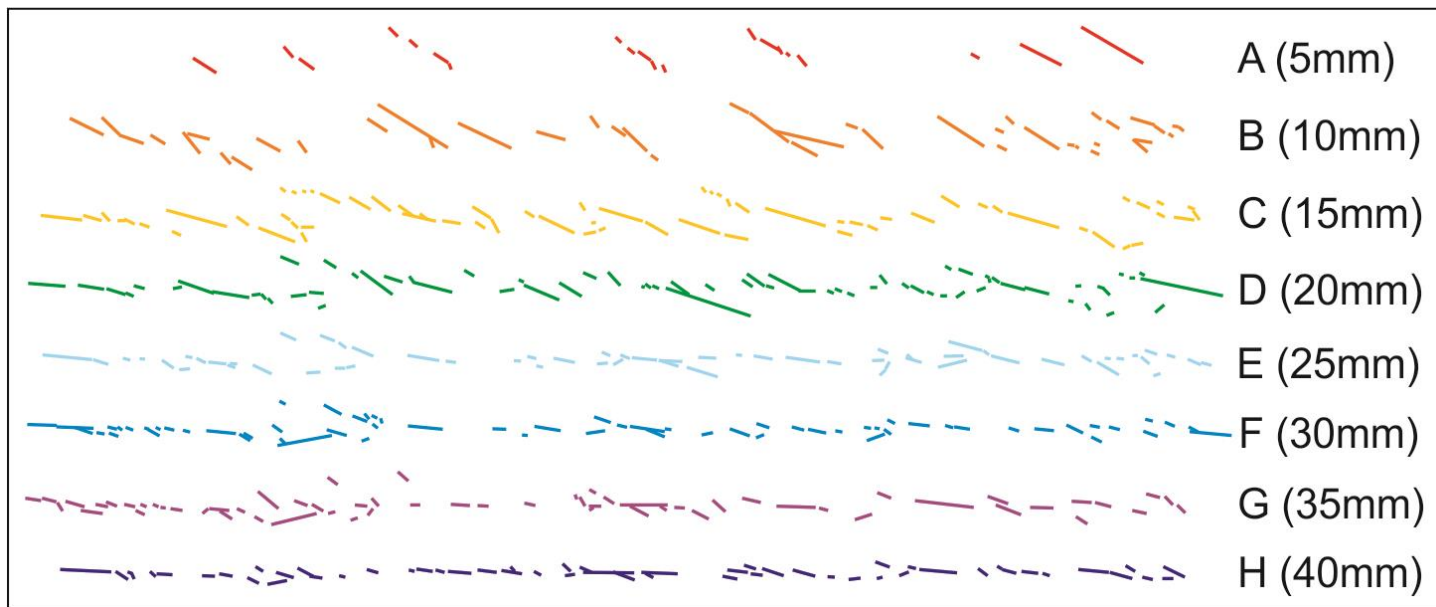
Multi layer model best replicates m-scale surface rupture morphology
Surface complexities created with simple, planar uniformly dipping basement fault beneath layered 'strata'

- Need to create a wide deformation zone but with discrete fractures

Comparative Deformation Zone Width (Distributed Shear) vs. Displacement

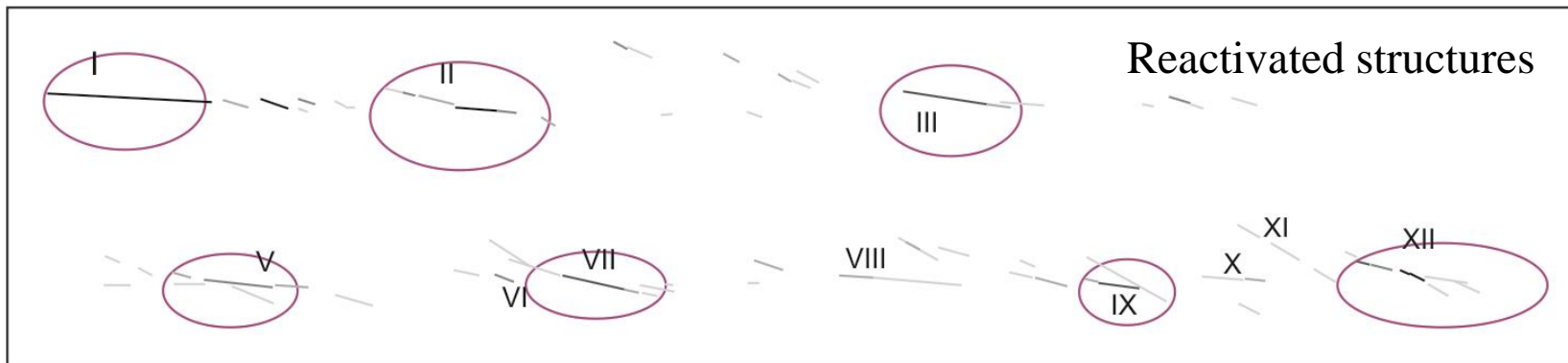
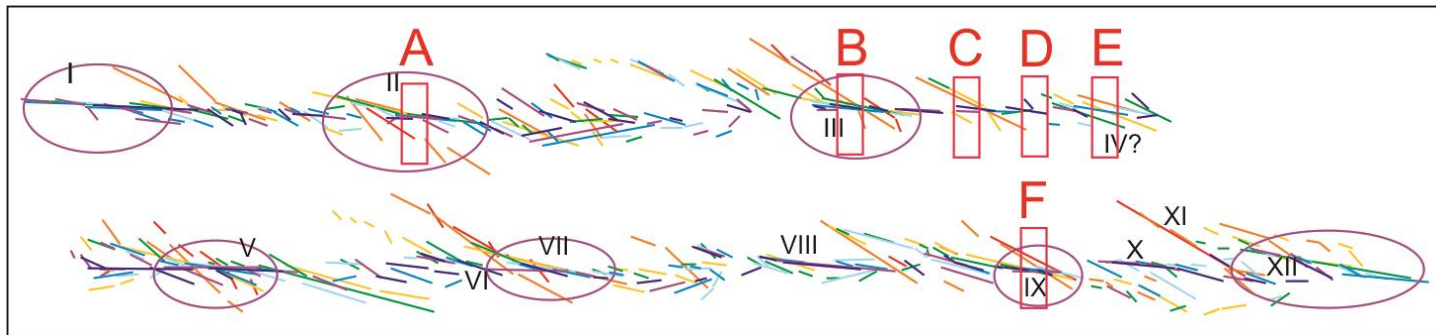


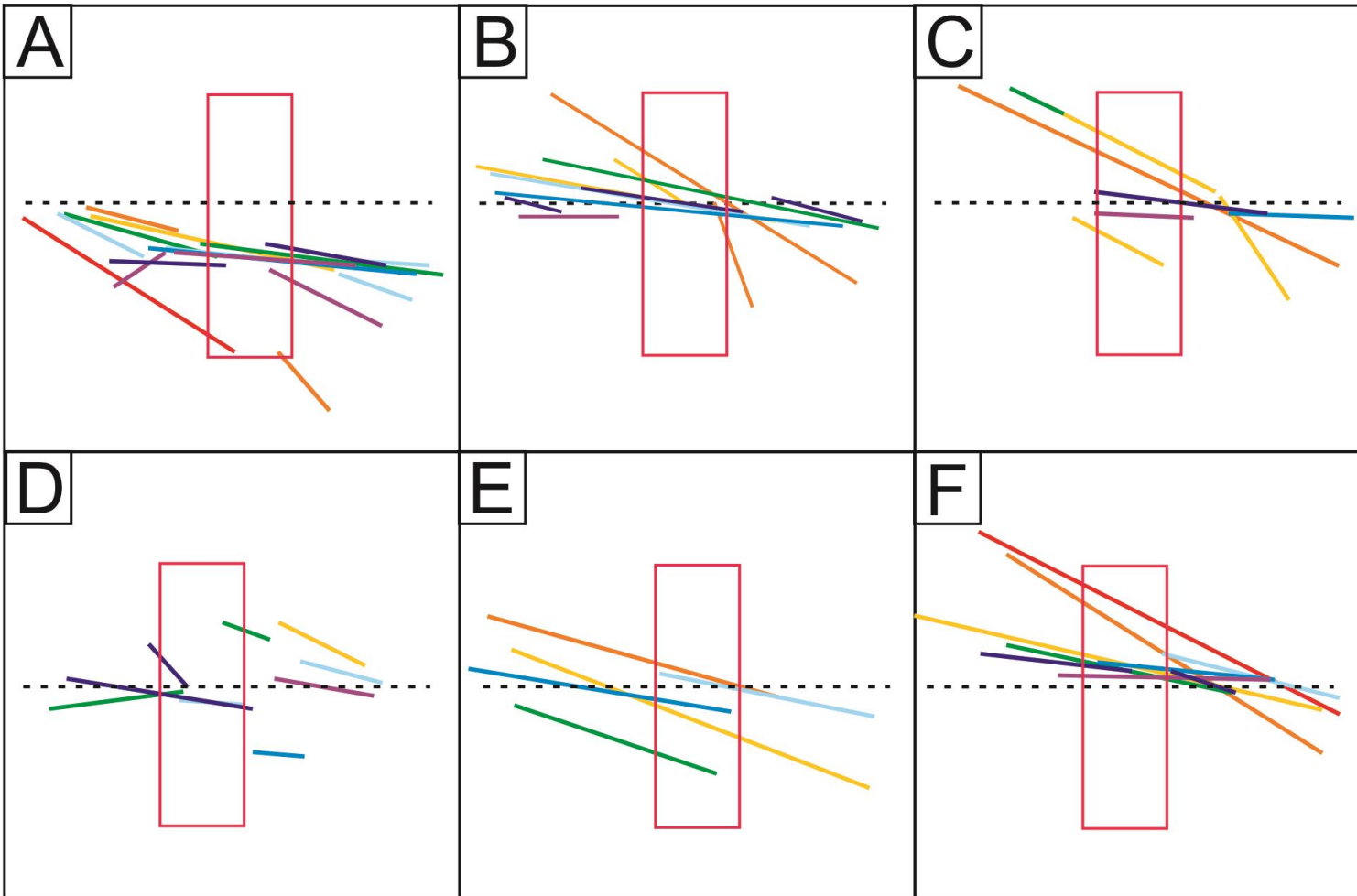
- Presence of granular sand increases shear zone width
- Presence of overlying cohesive layer concentrates distributed strain onto discrete fractures
- Best fit for surface rupture characteristics, but is this supported by subsurface geology?



•Successive
'active fracture'
mapping at
successive
strain
increments

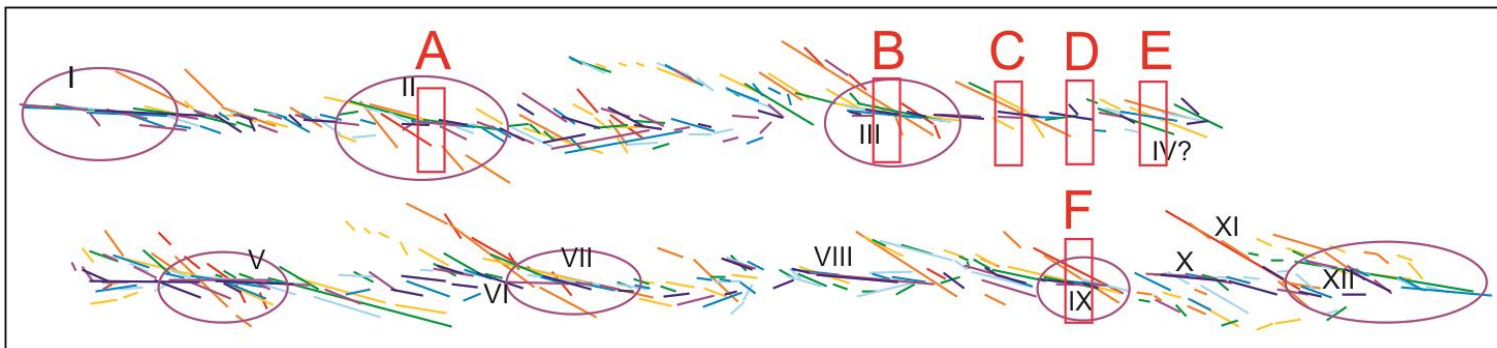
•Which
structures are
most
reactivated and
thus best
targets for
trenching?



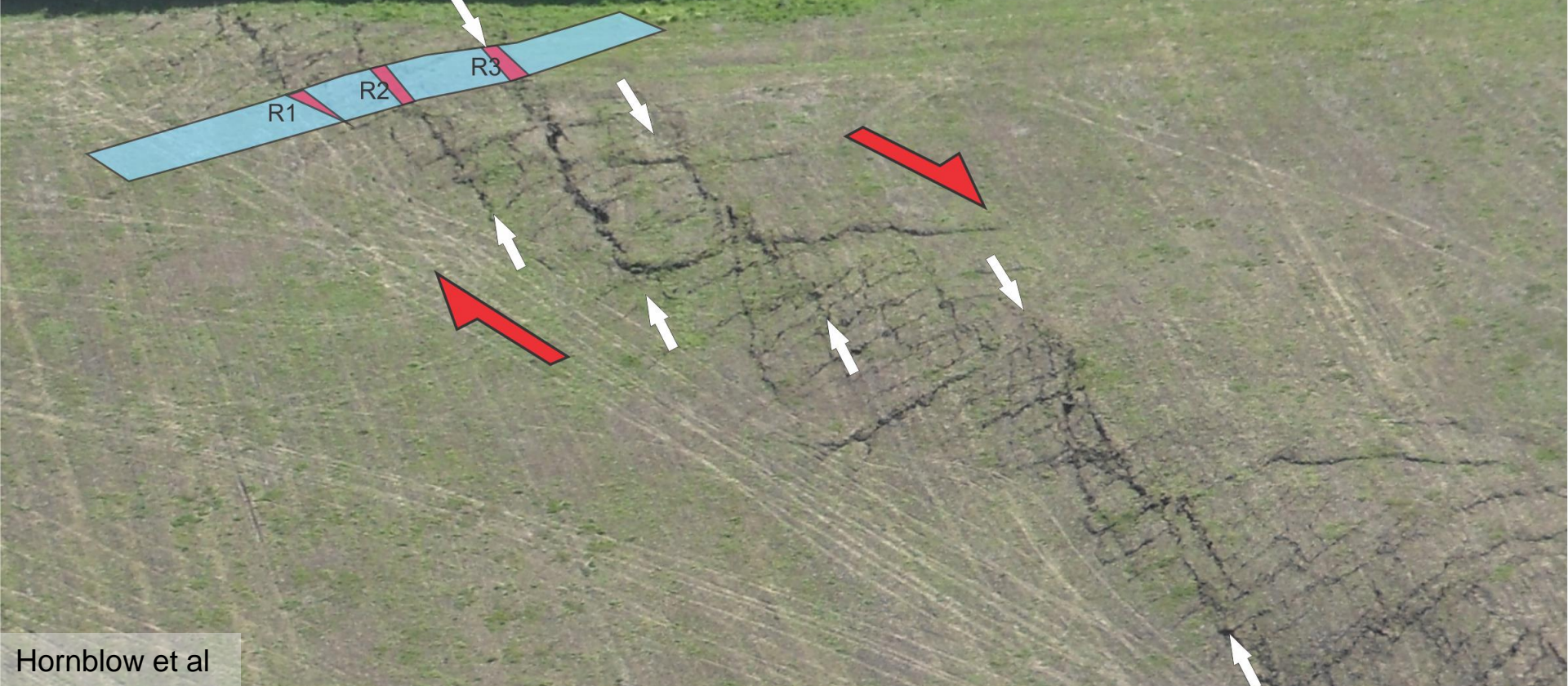
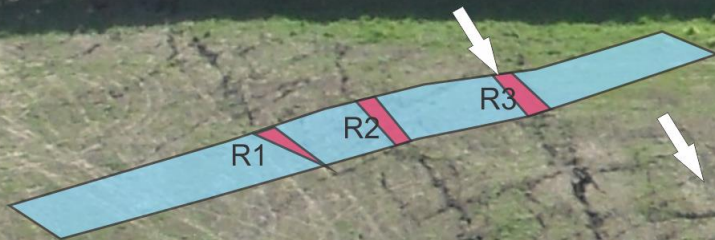


•Synthetic trenches – Low angle (Riedel) fractures best targets, but complex relationships possible

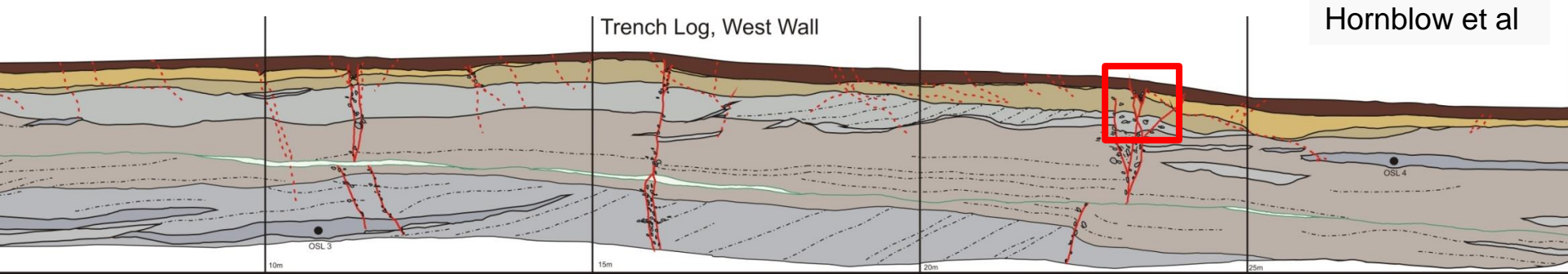
•Supported by trenching?



Greendale Fault paleoseismology project

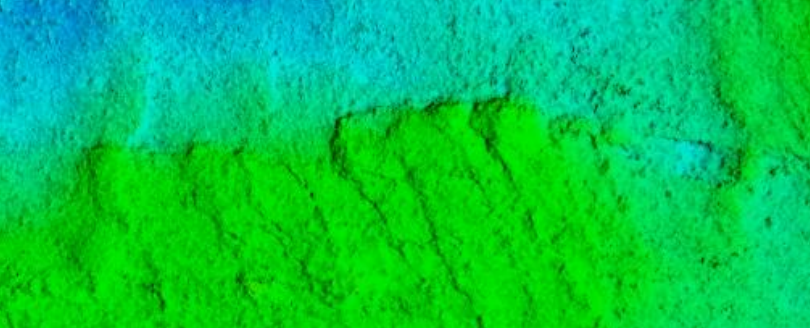


Subsurface mapping and GPR surveying of faulted sediments



- Many surface fractures terminate in uppermost 30-50cm (pedogenesis and loess filled channels increases cohesivity and promote fracturing)
- Thoroughgoing R fractures penetrate deeply and appeared to show more subsurface than surface displacement

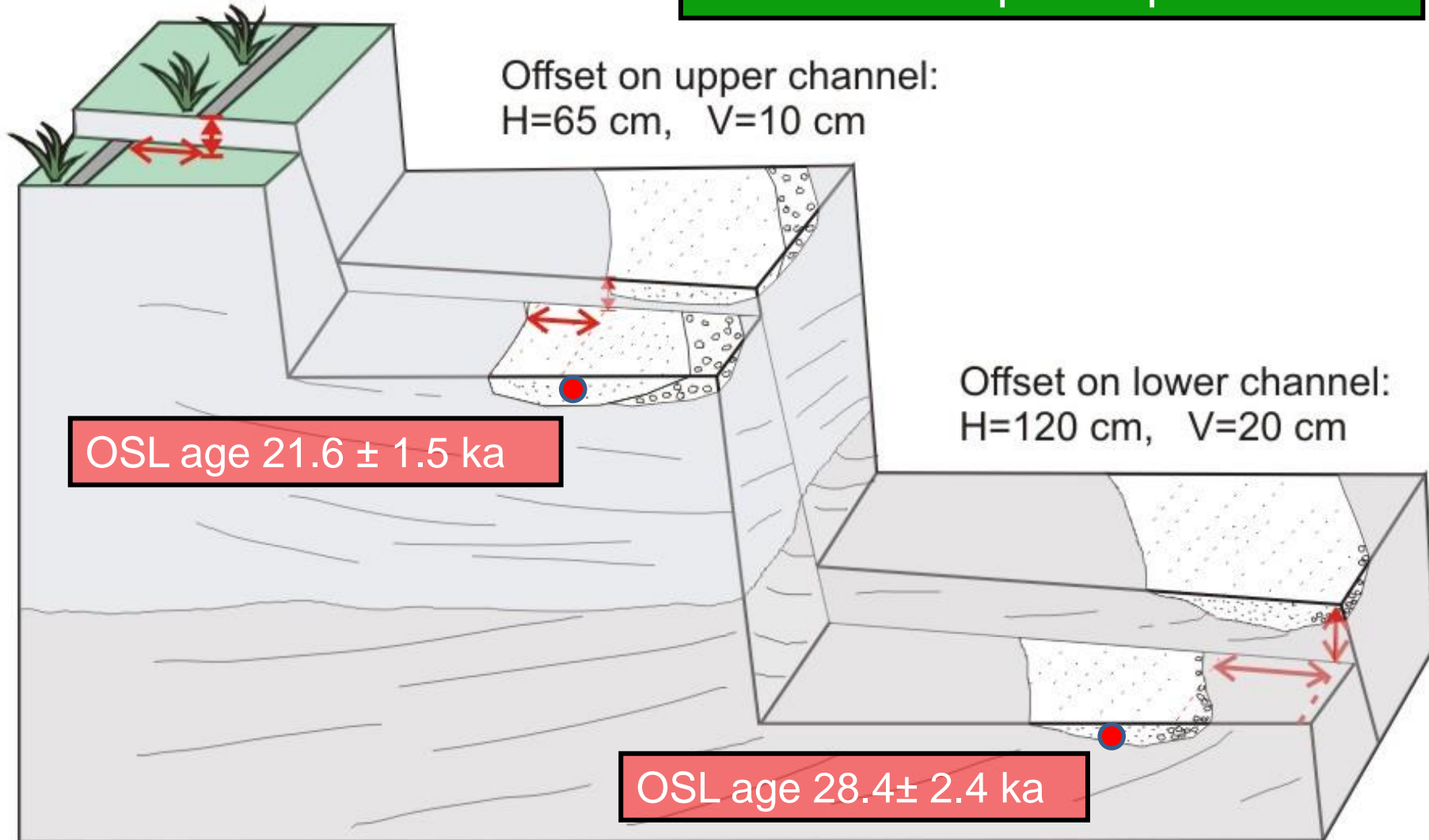




Digging laterally along fault to expose paleochannel cross-sections and measure piercing points (channel facies and margins)

**The penultimate earthquake:
Between ~22 and ~28 ka
Consistent slip-at-a-point**

2010 offset measured along structure
on surface $H = 60 \pm 10$ cm



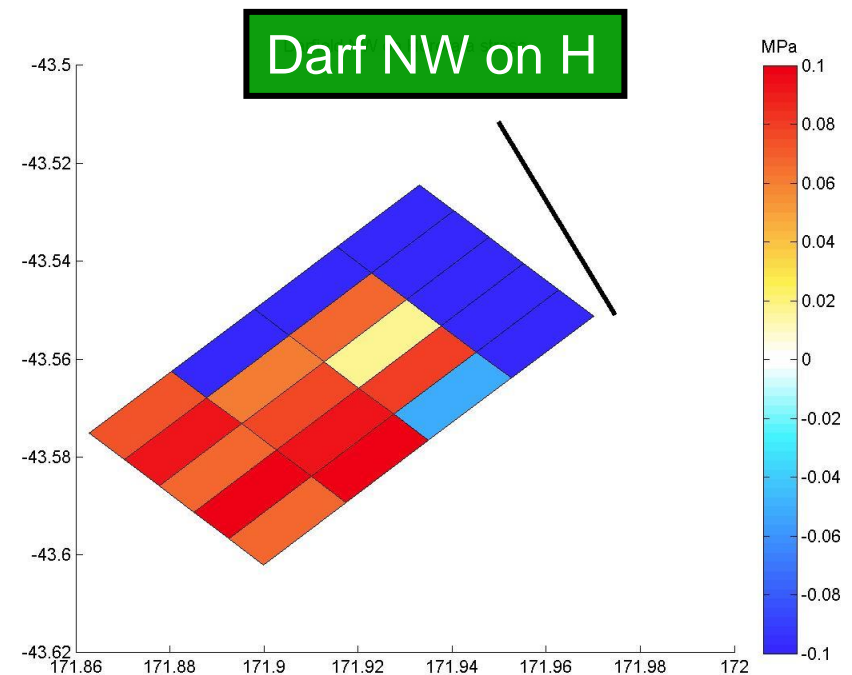
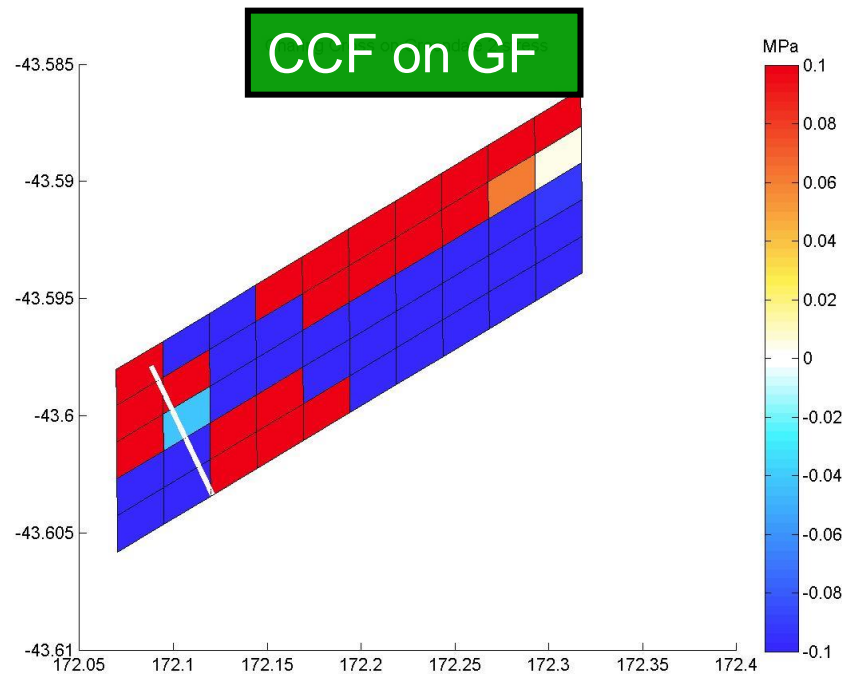
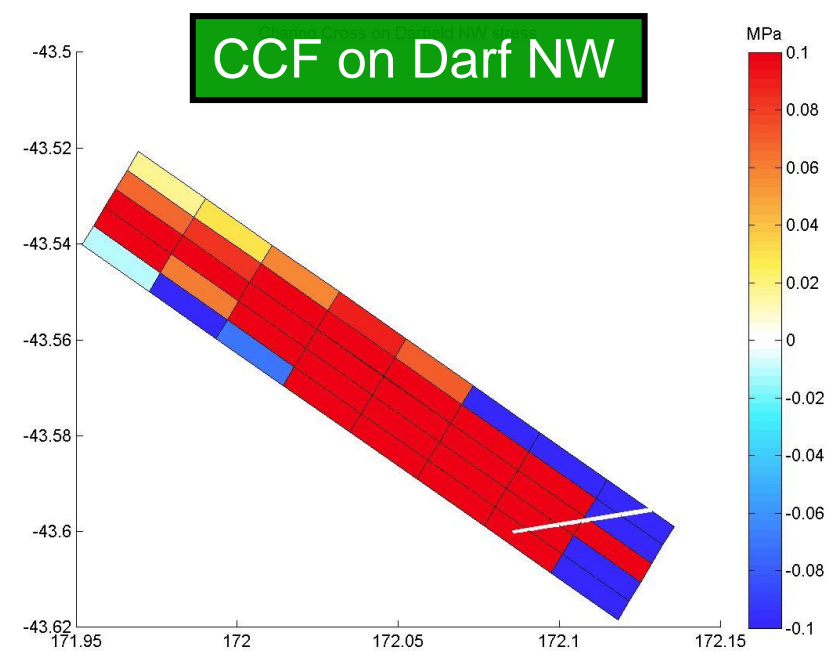
An aerial photograph of a dry lake bed, showing a complex network of cracks and a winding road. The cracks are dark and irregular, forming a web-like pattern across the light-colored, textured surface of the lake bed. A road, visible as a thin, light-colored line, winds through the landscape, particularly on the right side of the image. The overall scene suggests a geological or environmental study of a dry lake bed.

From point measurements to complex rupture scenarios

*What happens if the dominoes
topple the other way?*

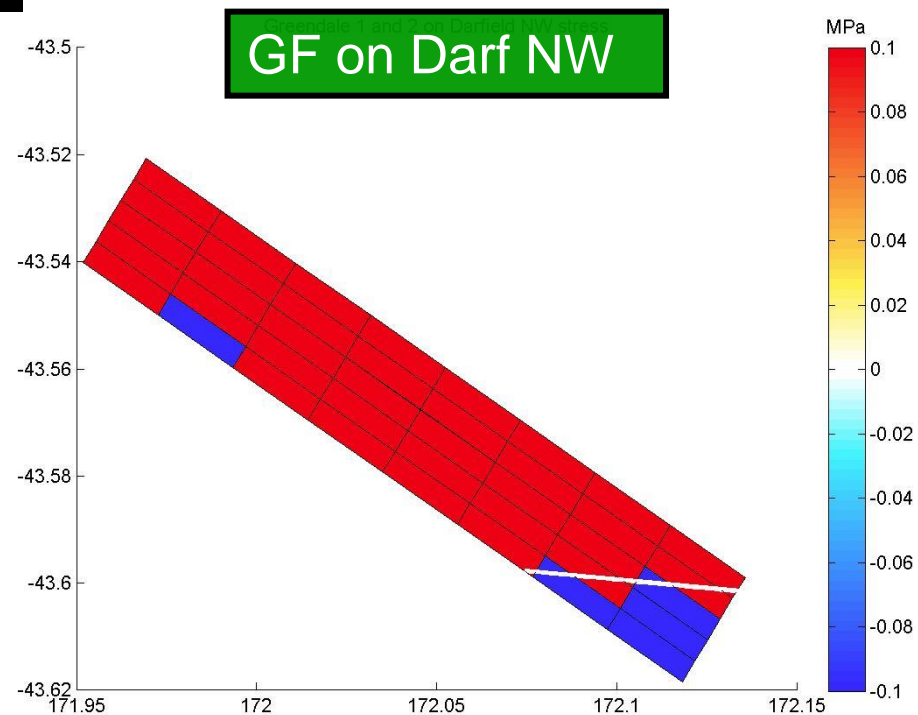
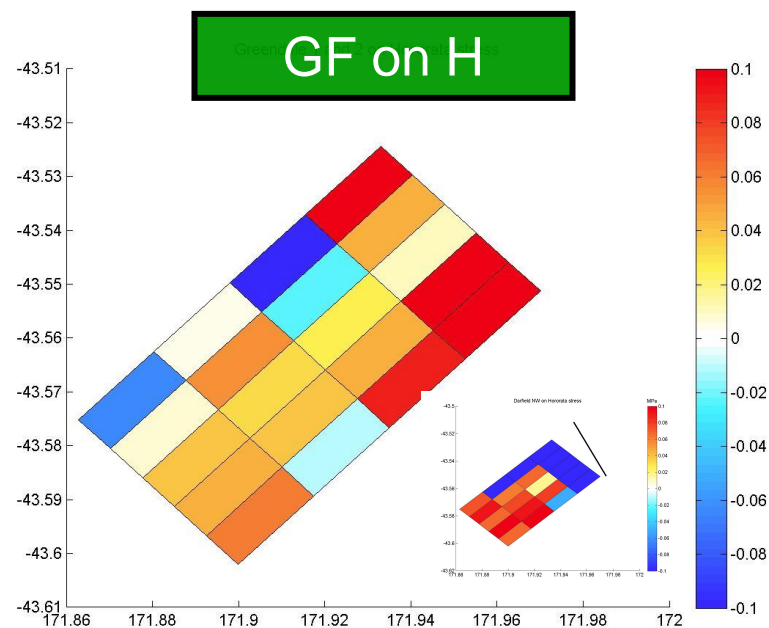
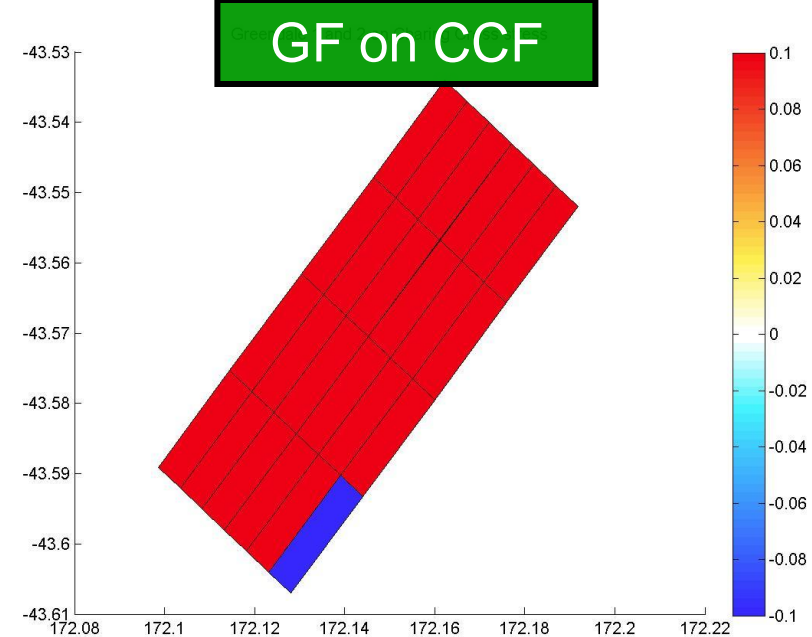


Coulomb 'static' stress evolution for rupture initiating on CCF



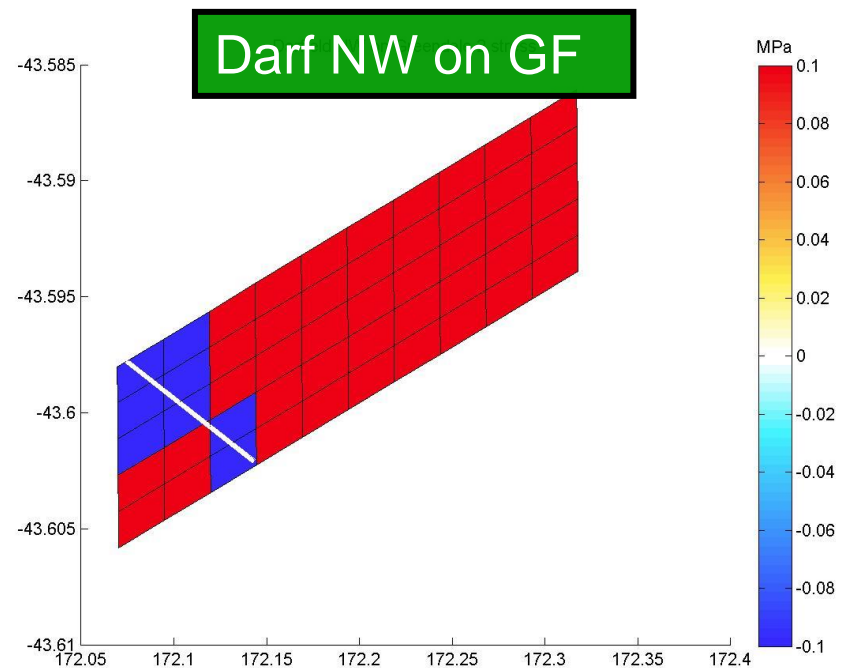
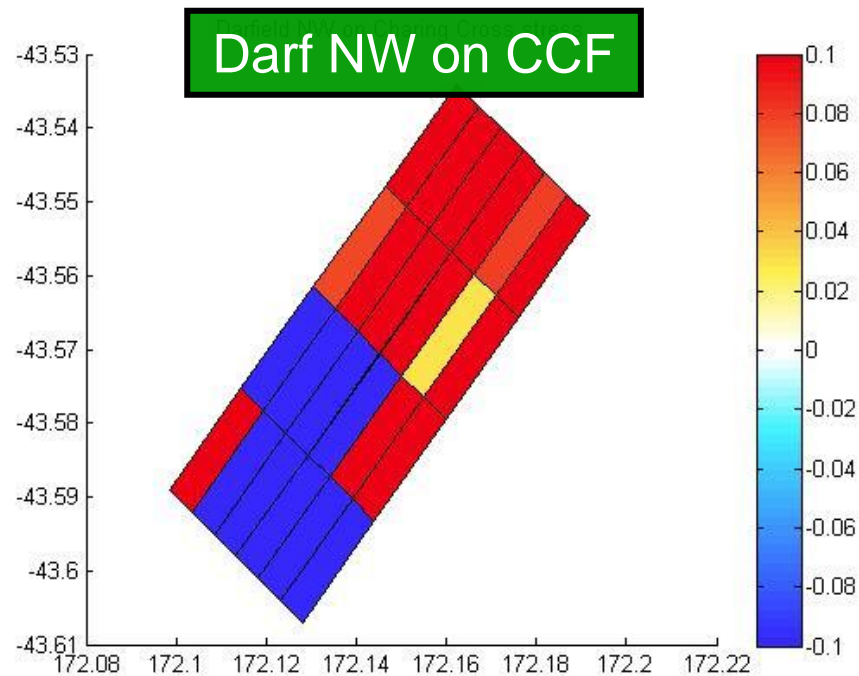
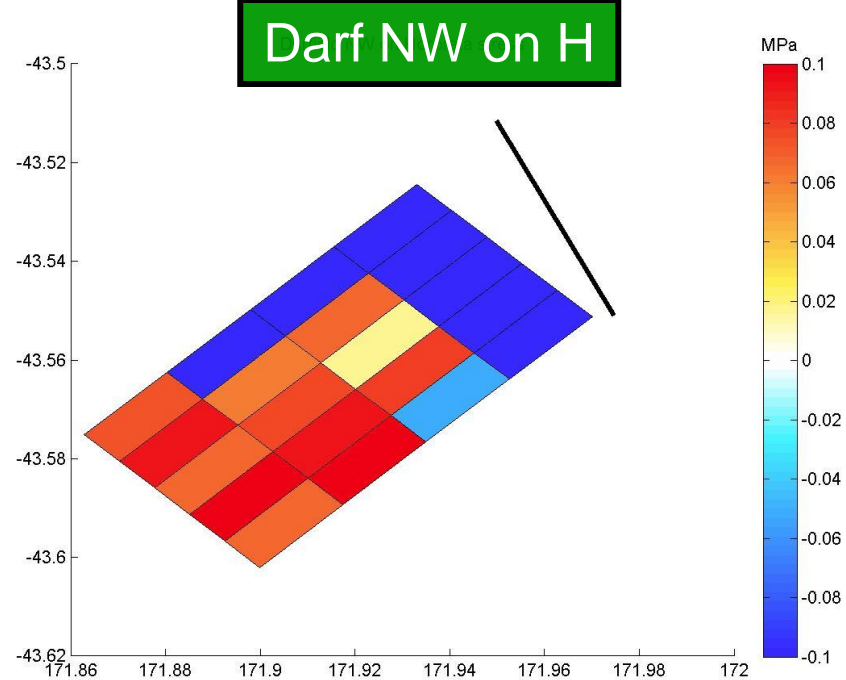


Coulomb 'static' stress evolution for rupture initiating on GF

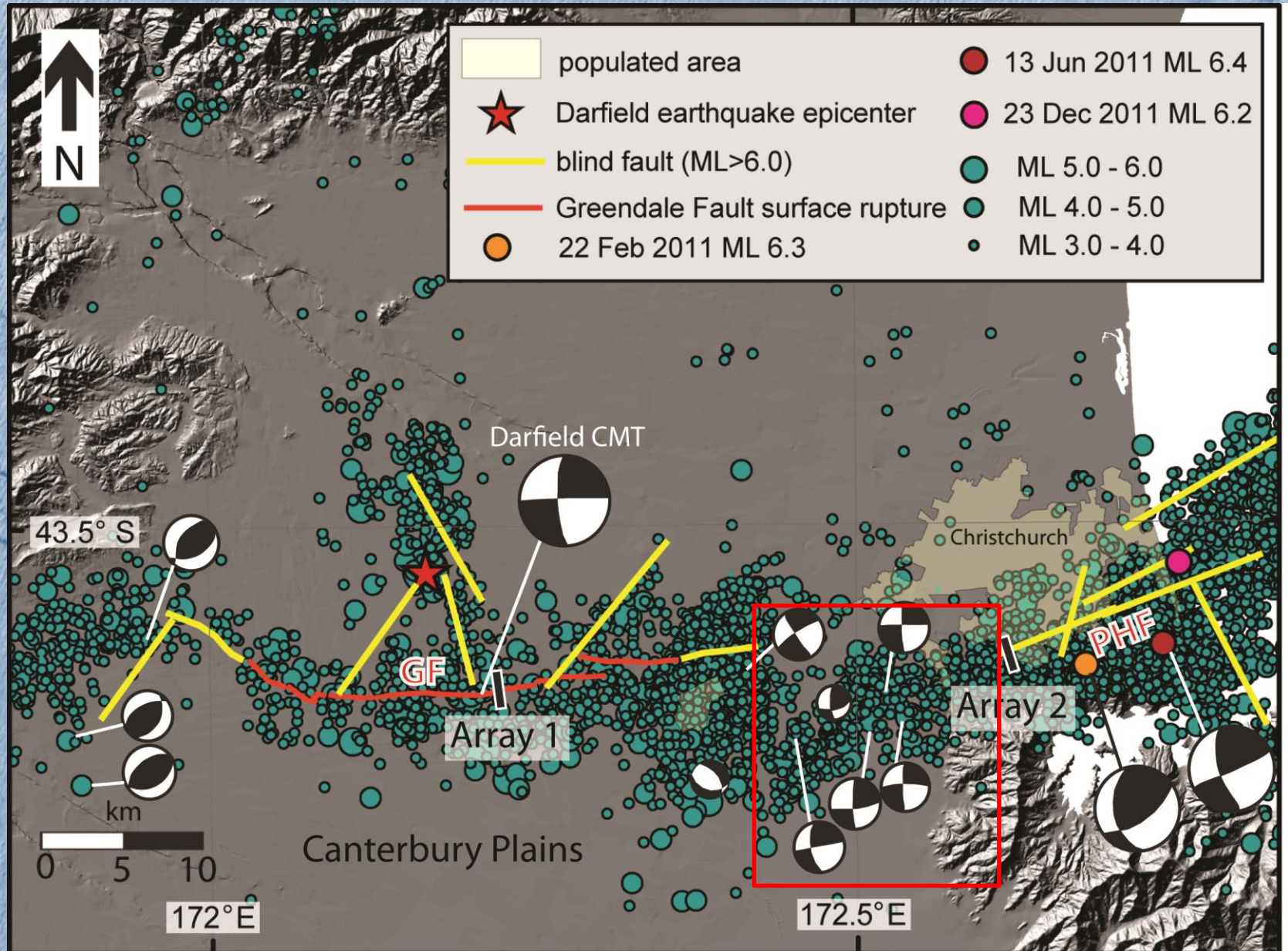


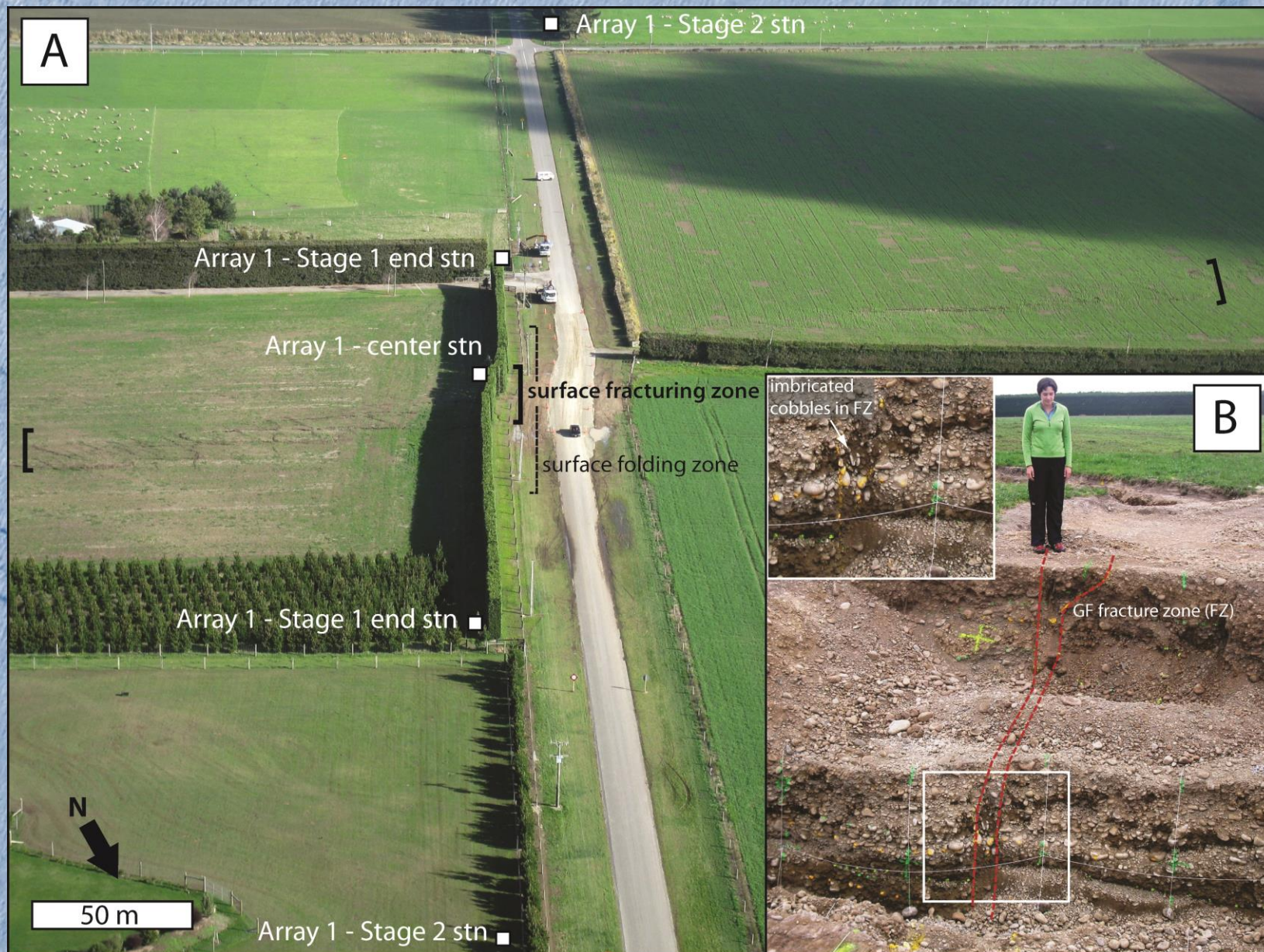


Coulomb 'static' stress evolution for rupture initiating on Darf NW



Other rupture scenarios (Mw max): Fault connectivity and rupture potential

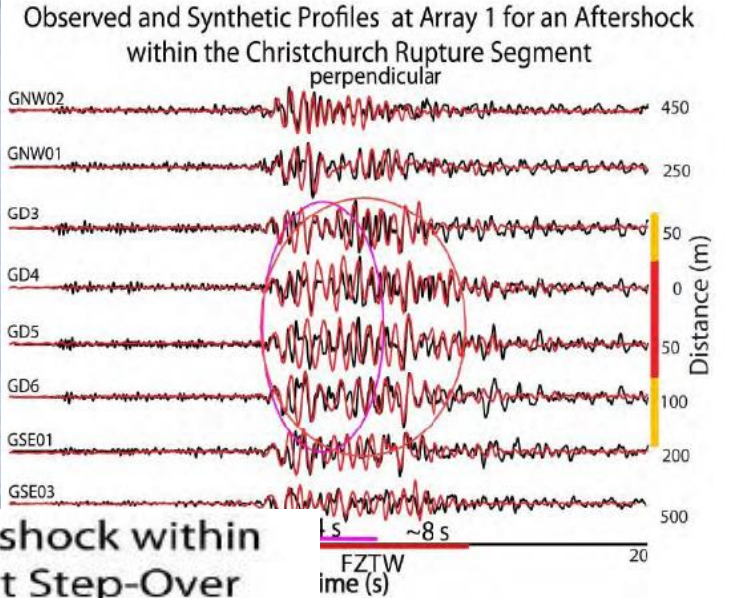




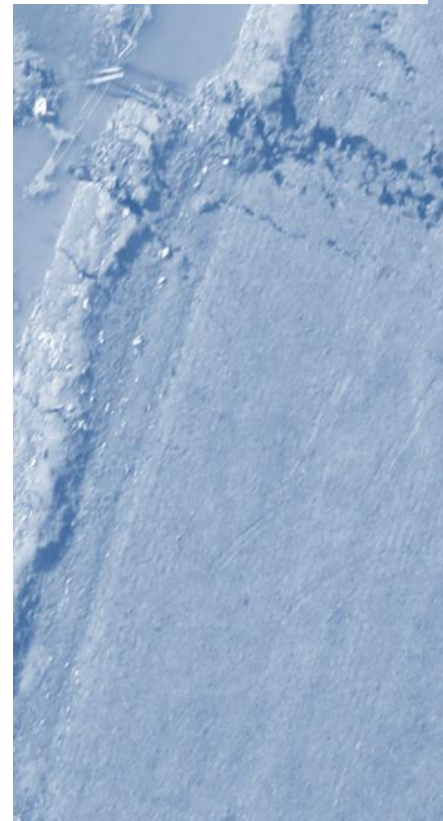
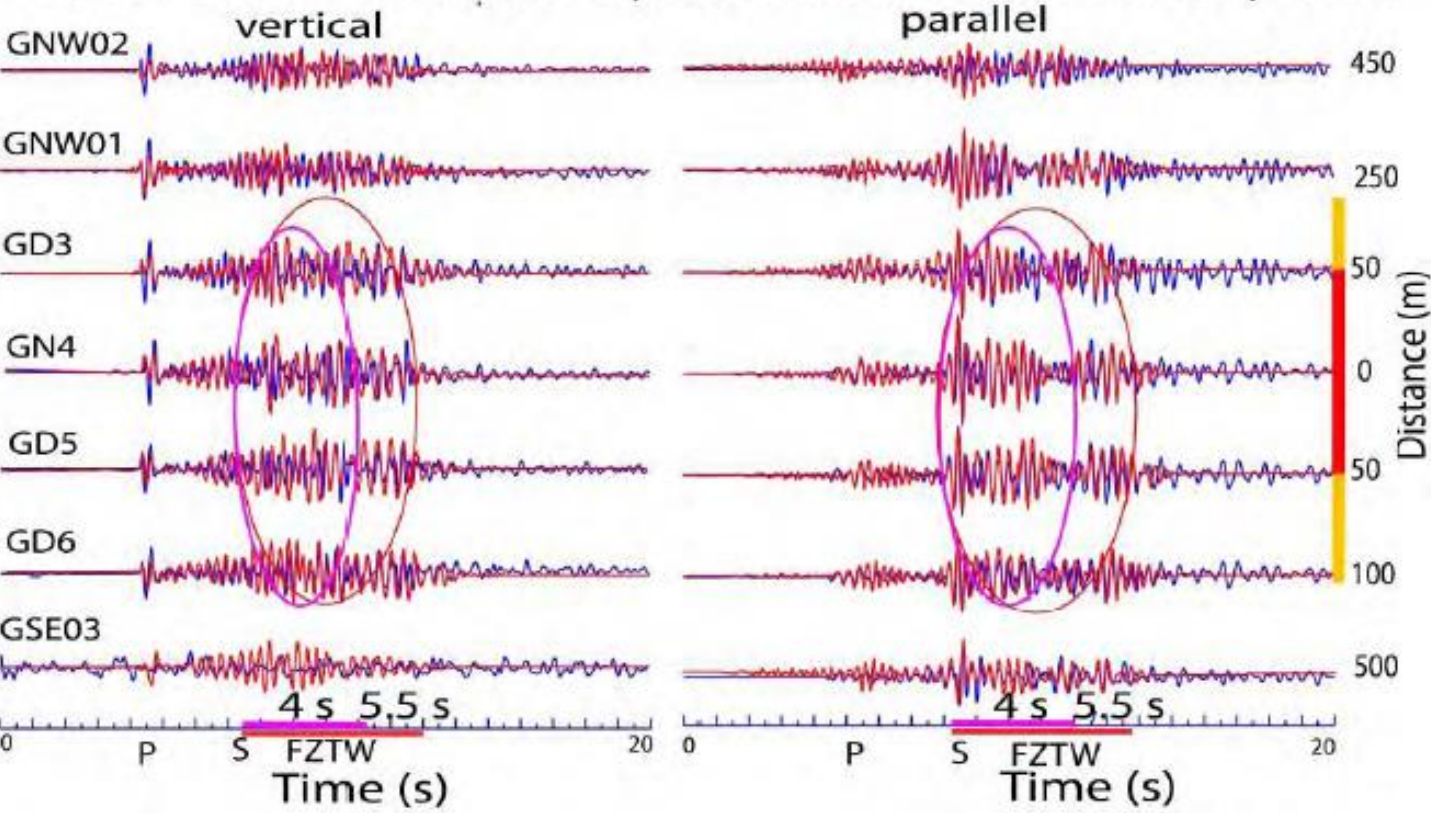
FZTW imaged on GF array for eq in 'The Gap' and on PHF

'Moderate connectivity' through gap via a complicated fracture mesh of small pre-existing faults and stress-aligned microcracks

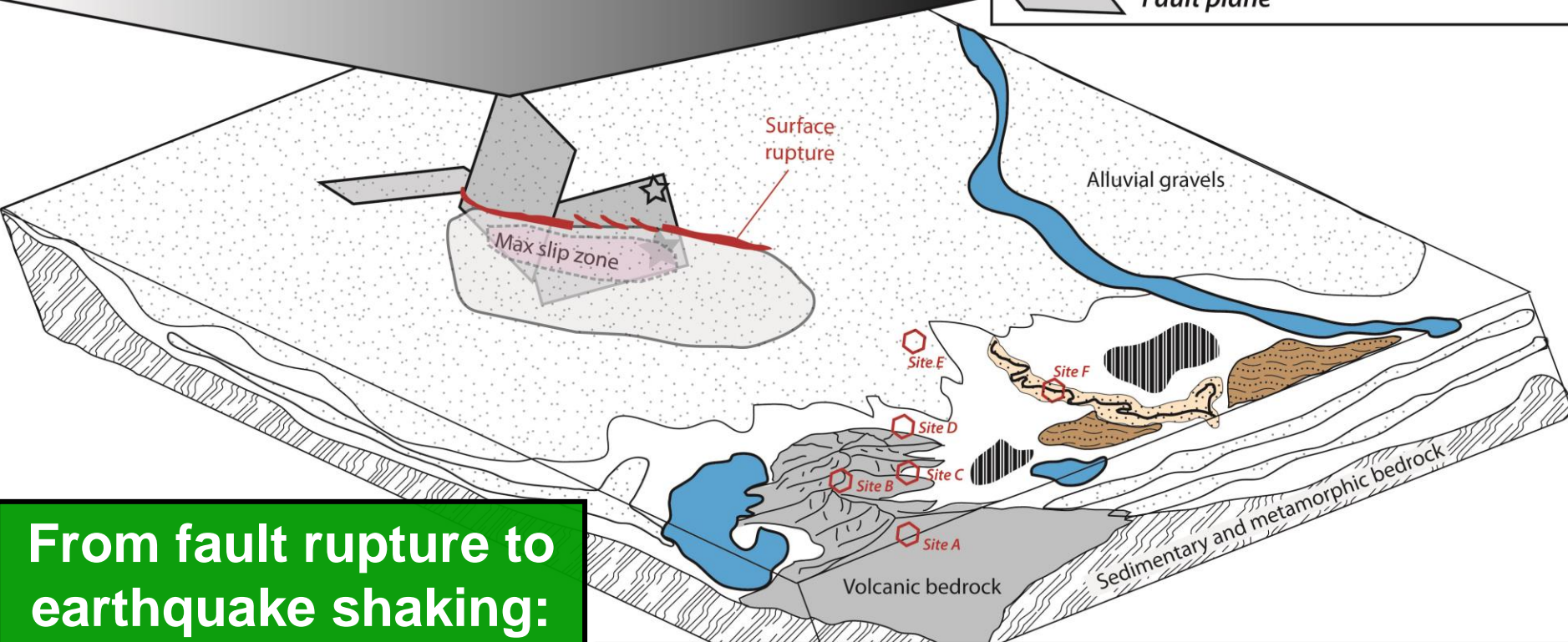
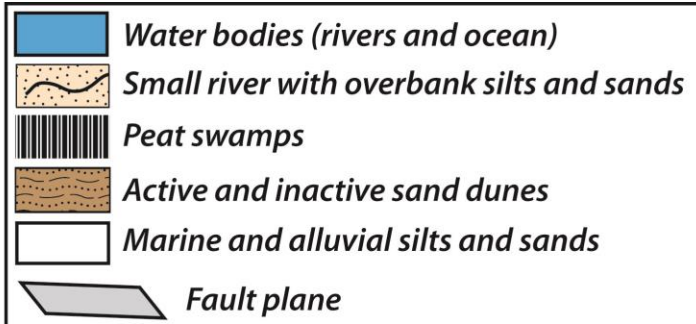
Improbable that these faults can rupture together (Mw 7.3 to 7.4) but this provides an example of an incipient system to compare to more mature faults: how do faults grow?



Observed and Synthetic Profiles at Array 1 for an Aftershock within East Extension of Darfield Rupture Zone beneath Fault Step-Over

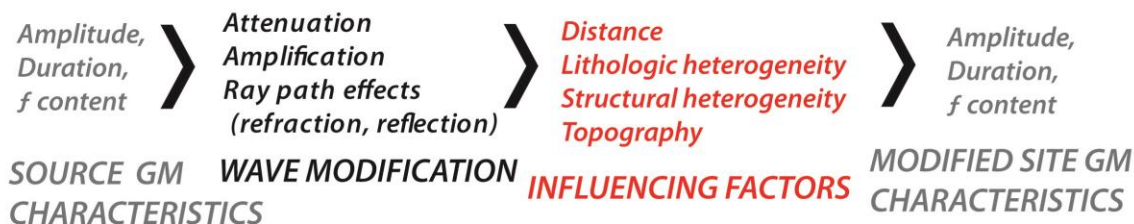


Relationships between seismic source and strong ground motion characteristics



From fault rupture to
 earthquake shaking:
 impacts and
 thresholds of
 geologic-geomorphic
 phenomena

Relationships between source and site strong ground motion characteristics



Each earthquake tells its own story:

Low frequency seismic amplification in the geologically-variable Christchurch “Jelly Bowl” and adjacent volcanic bedrock hills during the Feb 22 M_w 6.2 earthquake

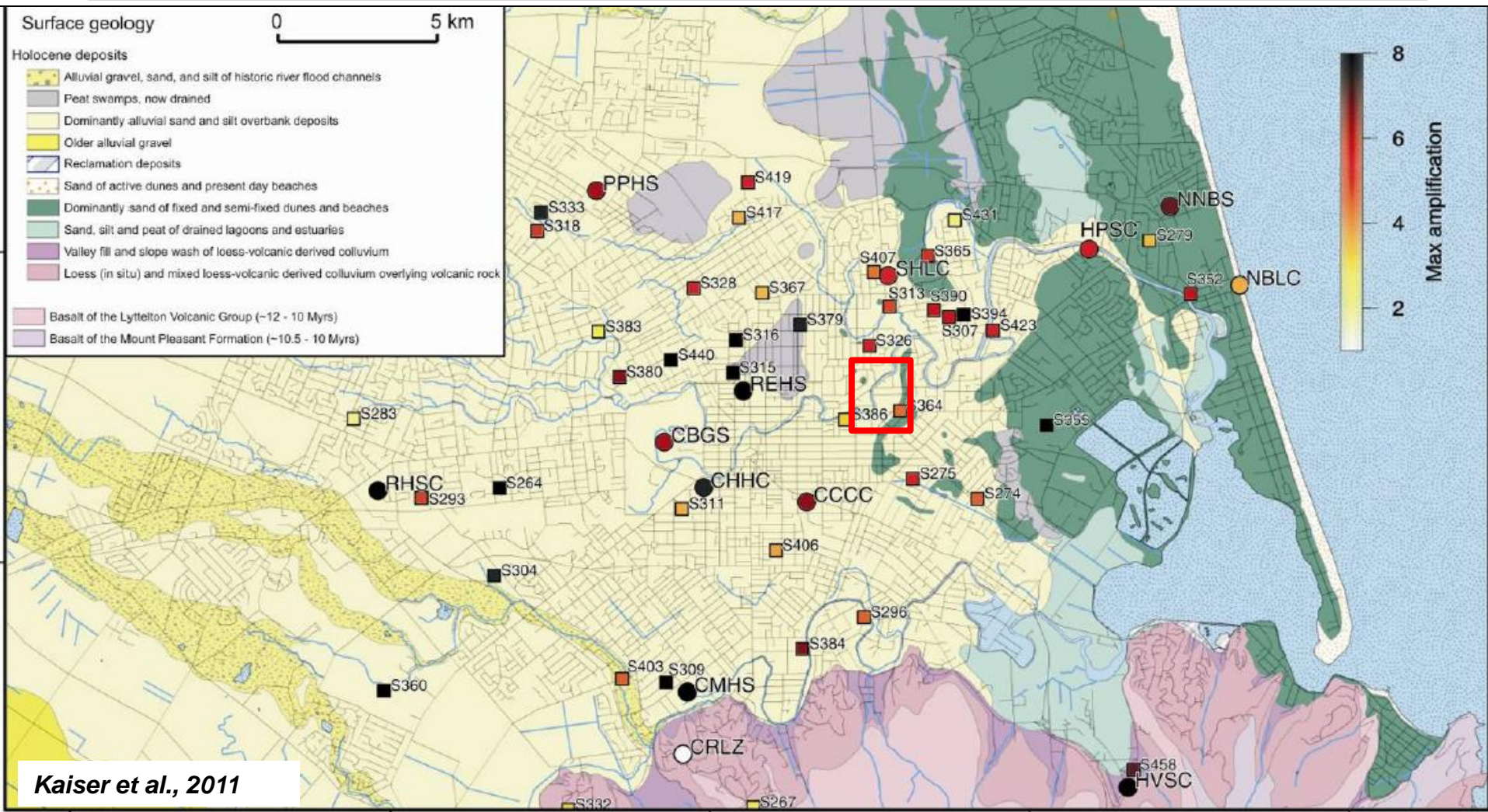
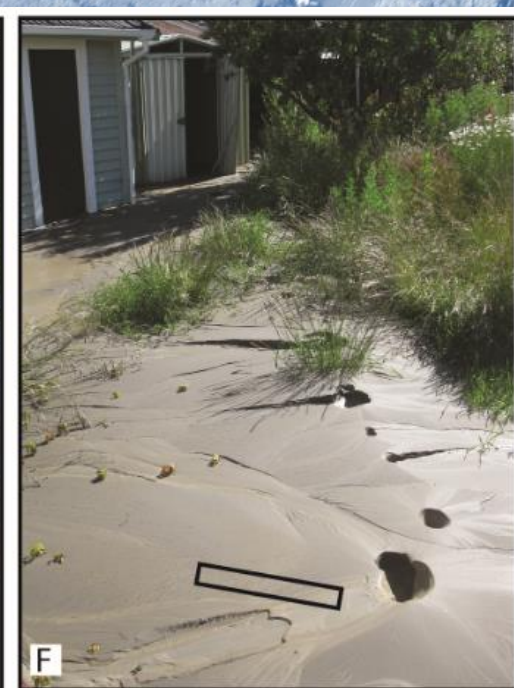
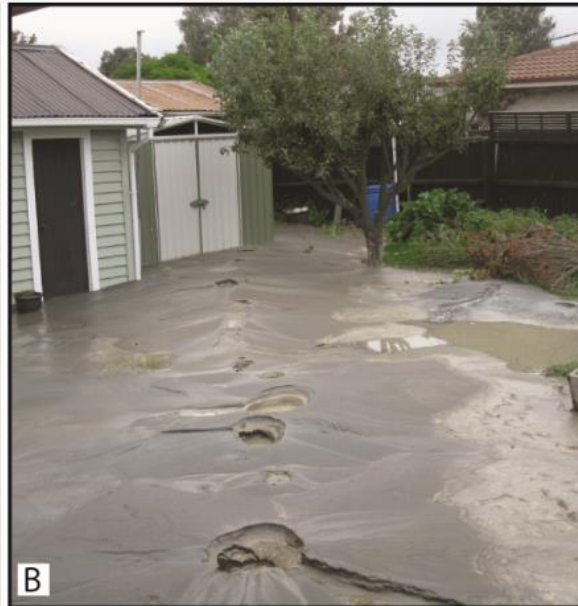
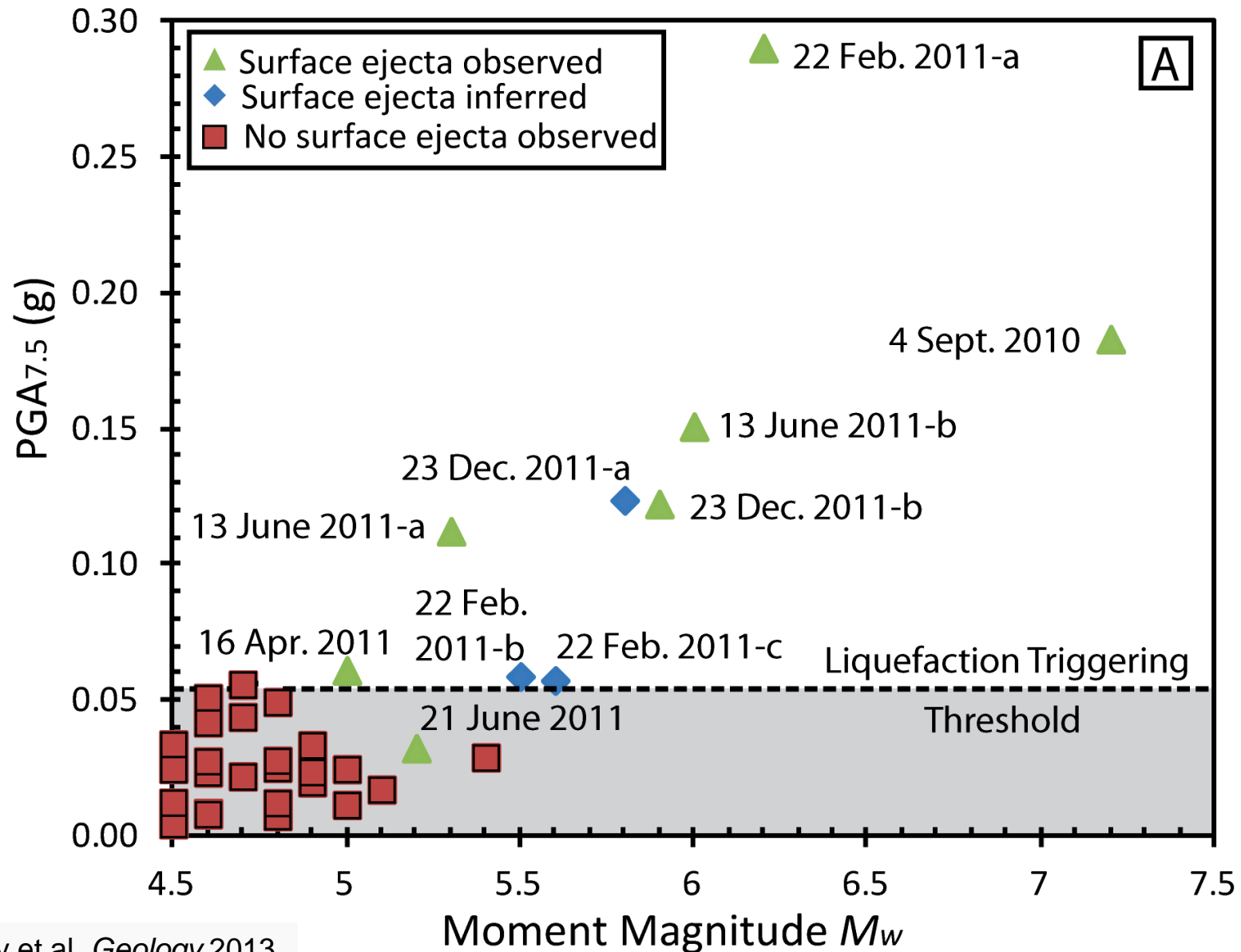


Figure 7. Maximum amplification in the 1 - 9 Hz frequency band derived from spectral ratio calculations at GeoNet stations (circles) and QCN stations (squares). Warmer station colours indicate higher amplifications relative to reference station CRLZ. Background map shows surface geology of the Christchurch area following Brown and Weeber 1992). Coordinates are New Zealand Map Grid given in metres.

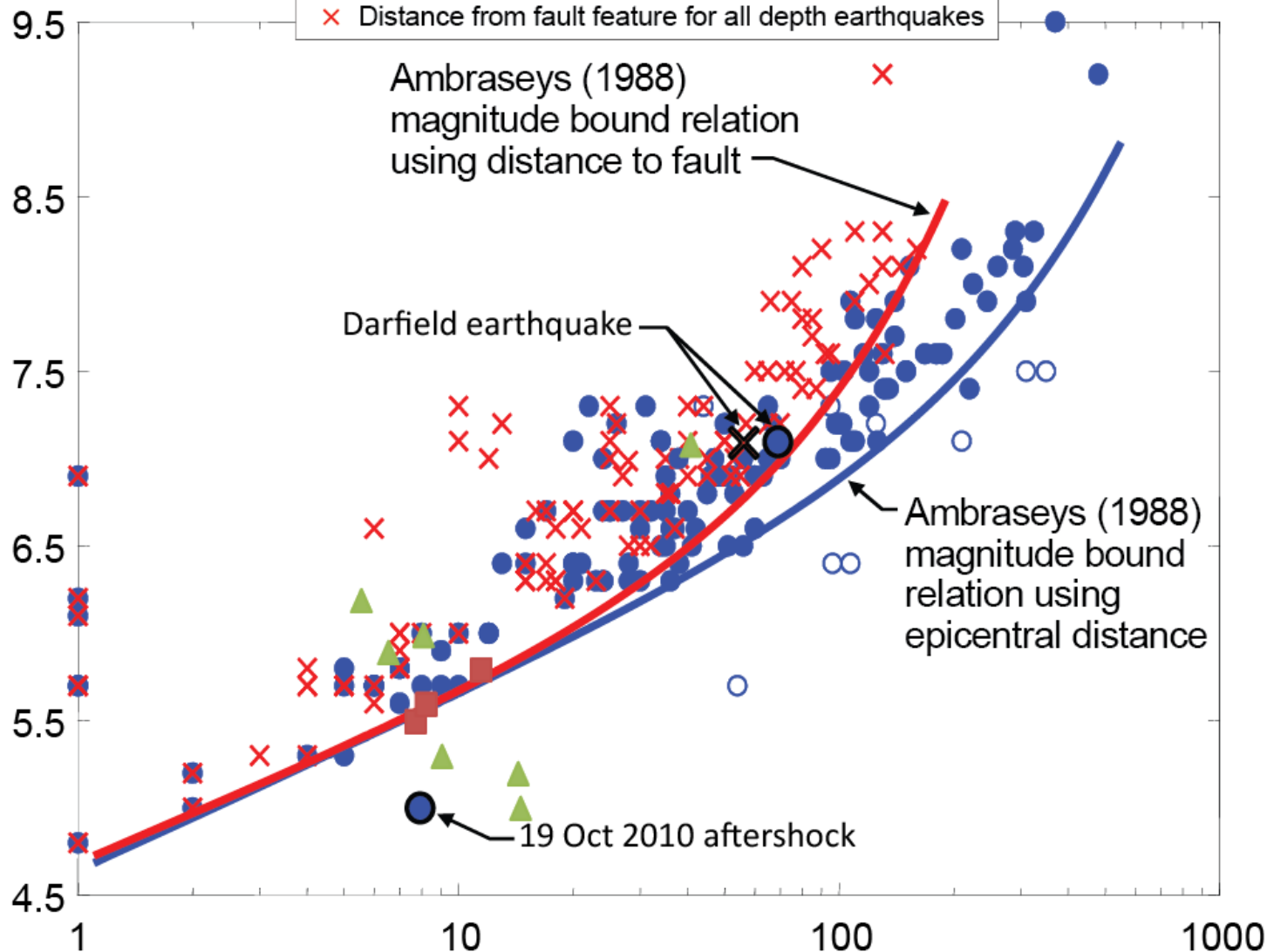
Recurrent liquefaction in Christchurch during the Canterbury earthquake sequence



At least 10 liquefaction episodes, some of which occurred at surprisingly low PGAs



Earthquake Moment Magnitude (M)



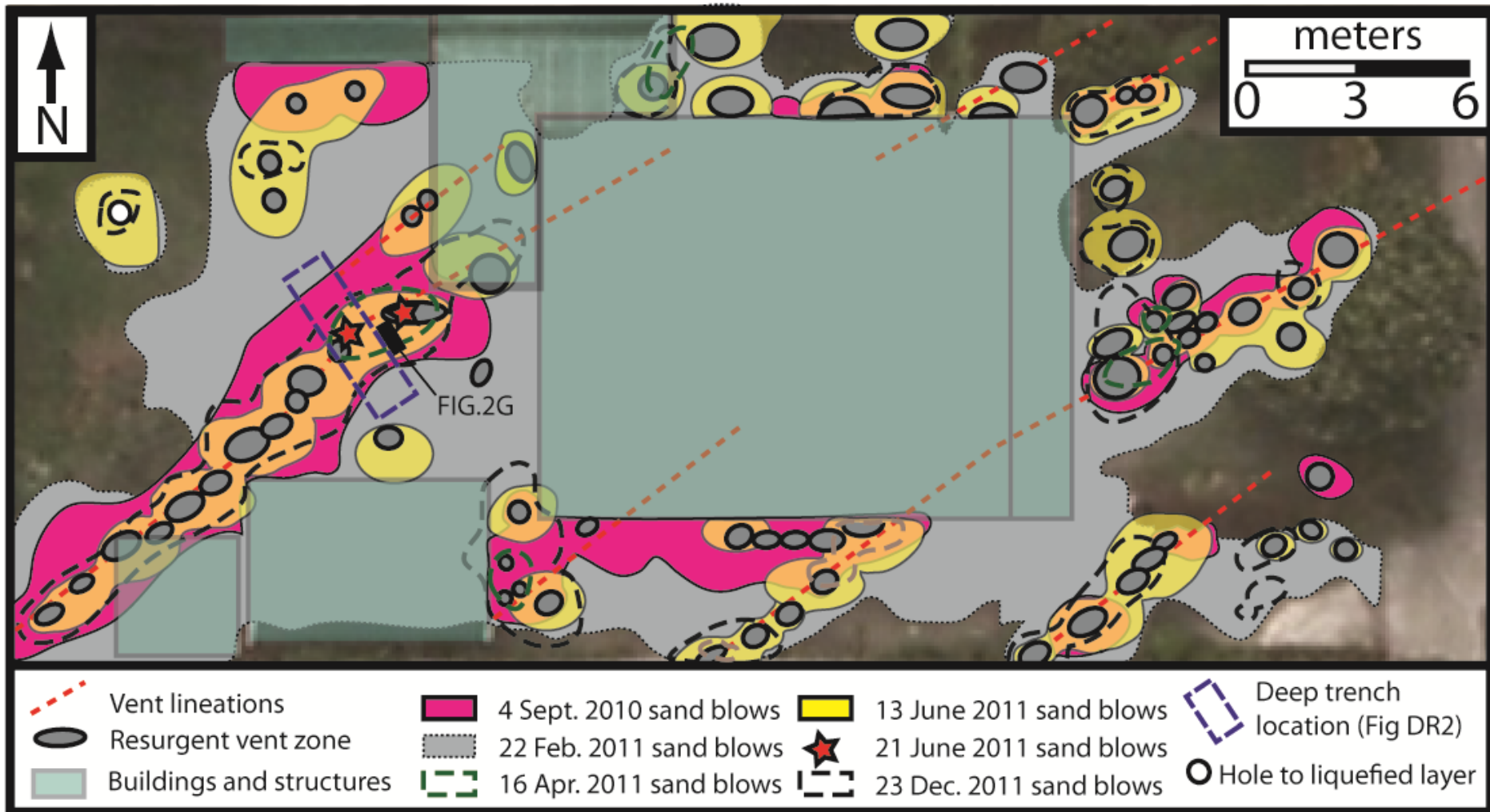
- Epicentral distance for shallow to intermediate depth earthquakes (focal depth < 50 km)
- Epicentral distance for intermediate to deep earthquakes (focal depth > 50 km)
- × Distance from fault feature for all depth earthquakes

Avonside study site

▲ Liquefaction features observed

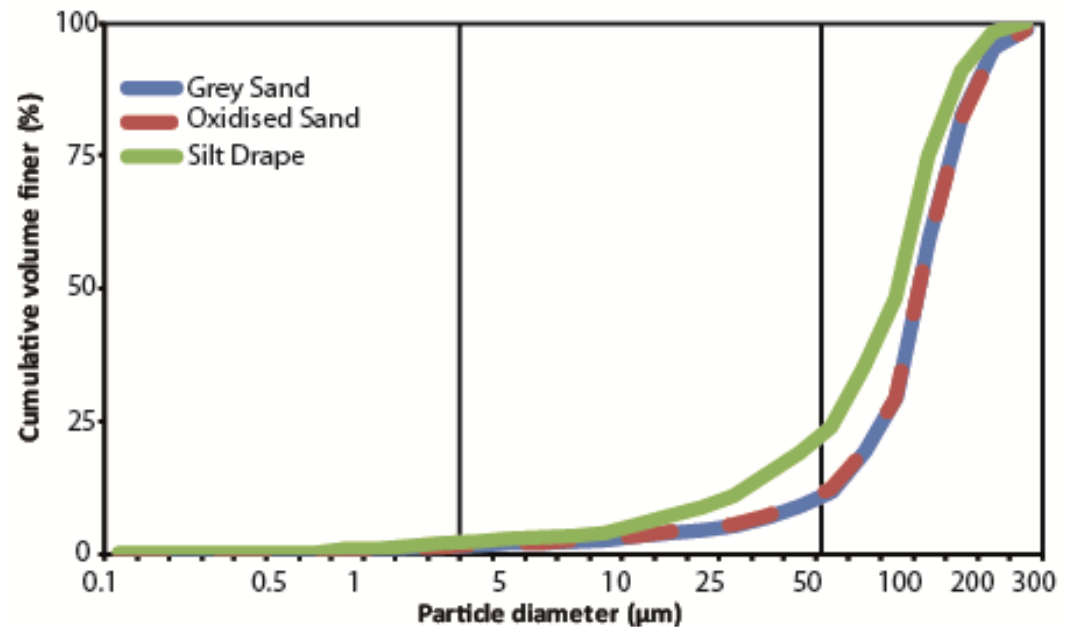
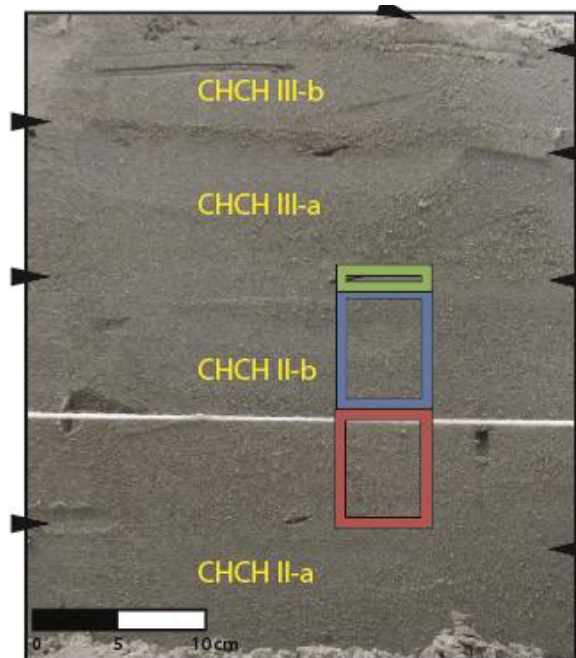
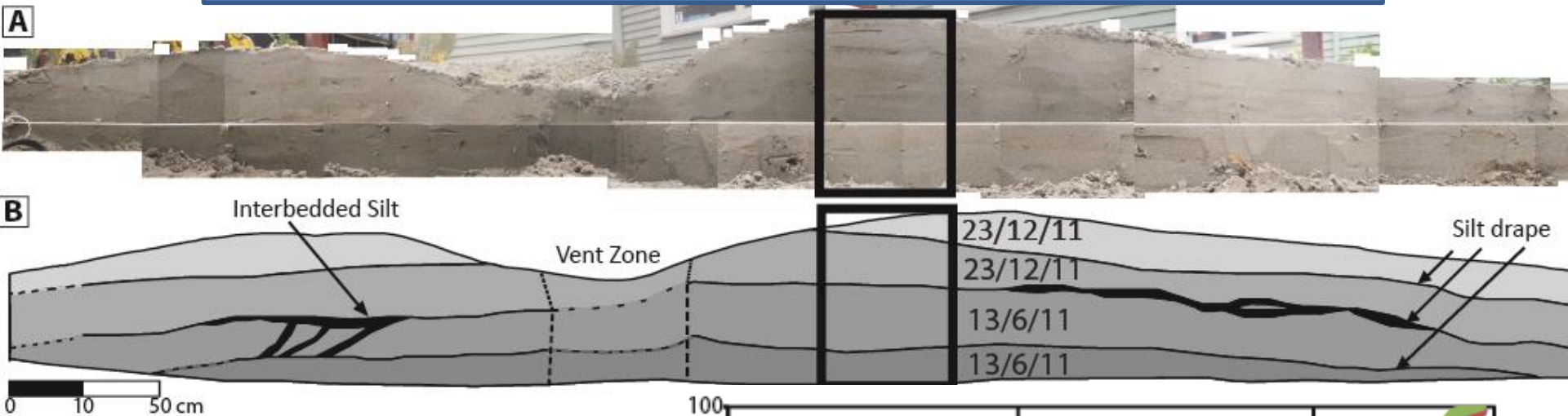
■ Liquefaction features inferred

Liquefaction sourced from the same vent structures related to lateral spreading – might prior events have done the same?



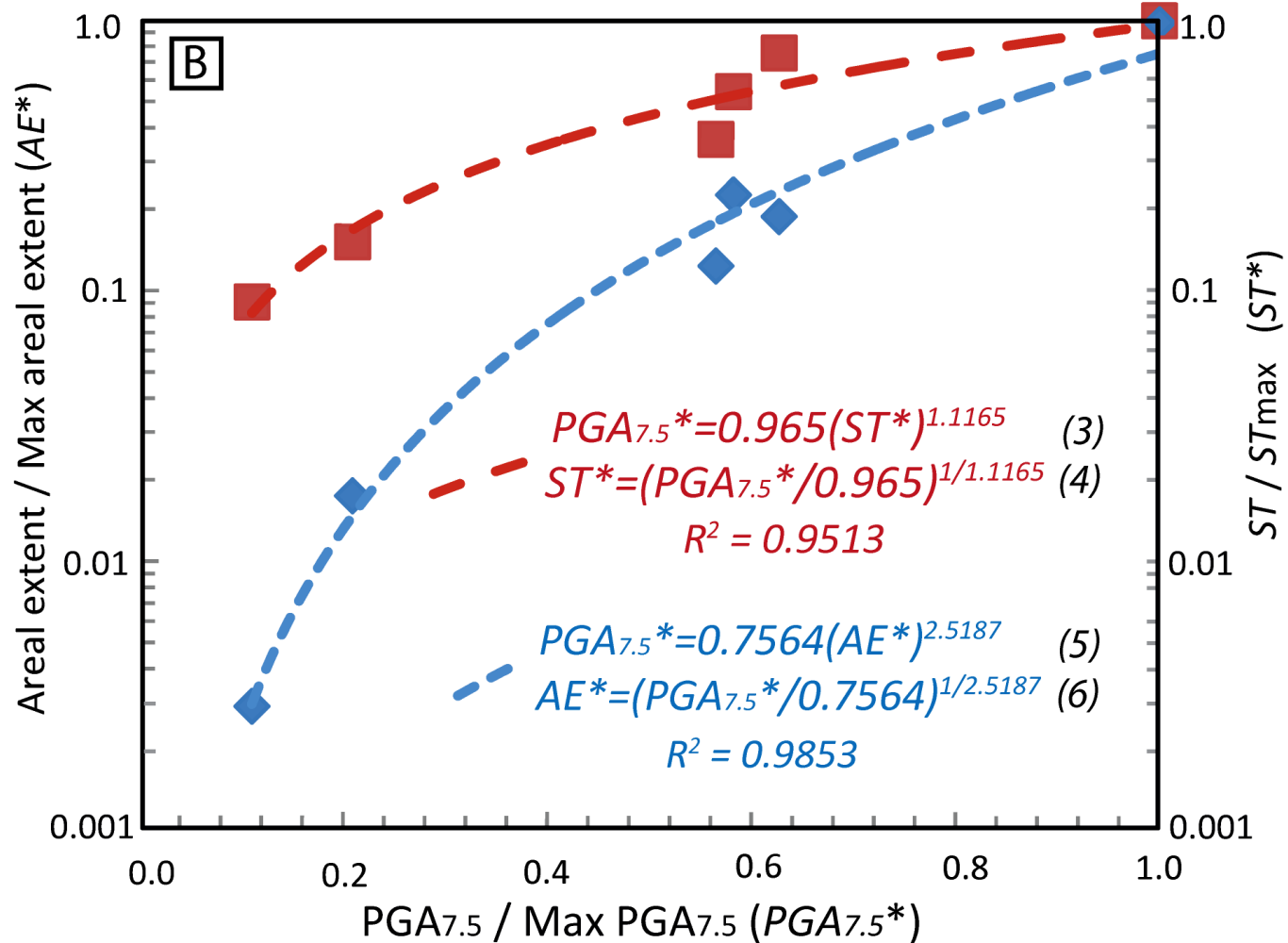
Would we recognize this in the geologic record?

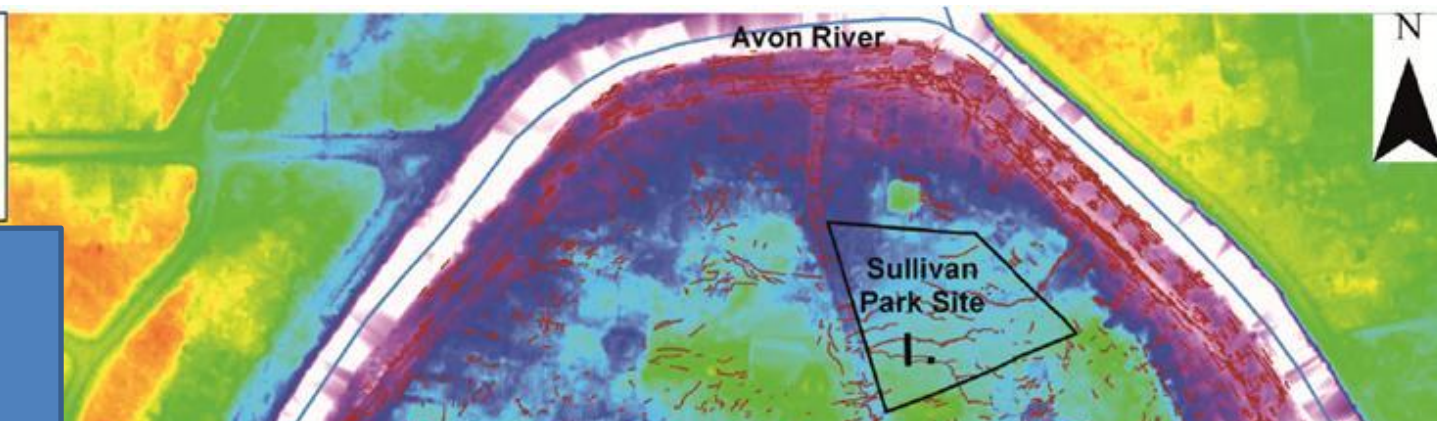
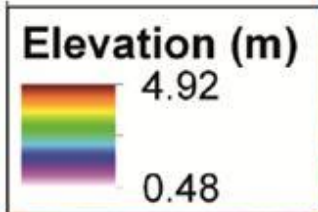
Surface ejecta trenching – develop relationships between earthquake characteristics and extent / thickness of sand blows



Clay	Very Fine	Fine	M	Coarse	V Fine	Fine
	Silt				Sand	
	4				62	

Normalize maximum thicknesses, areal extents and PGAs to maximum values: predict future sand blow characteristics from PGAs and interpret geologic record of paleoearthquakes





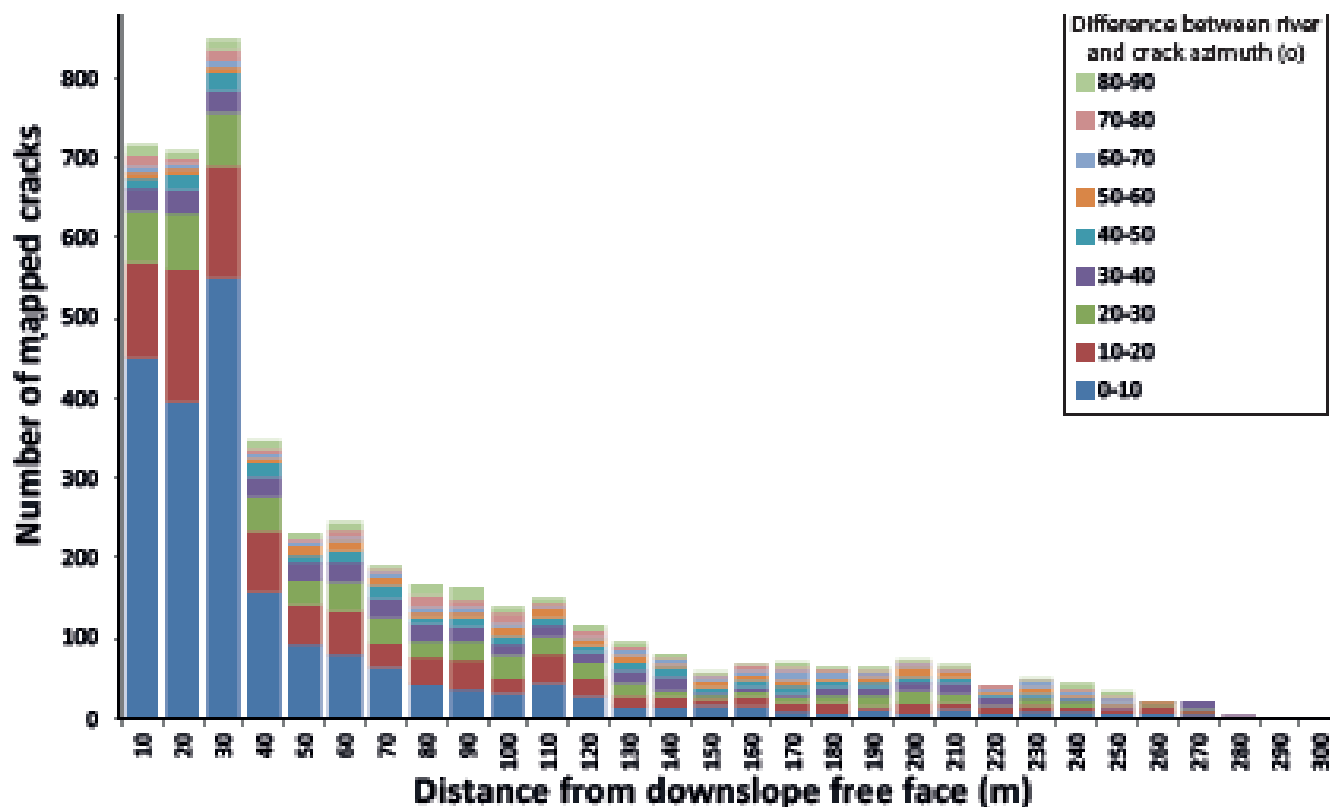
Where to look for paleoliquefaction?

Mapping of 4875 lateral spreading cracks

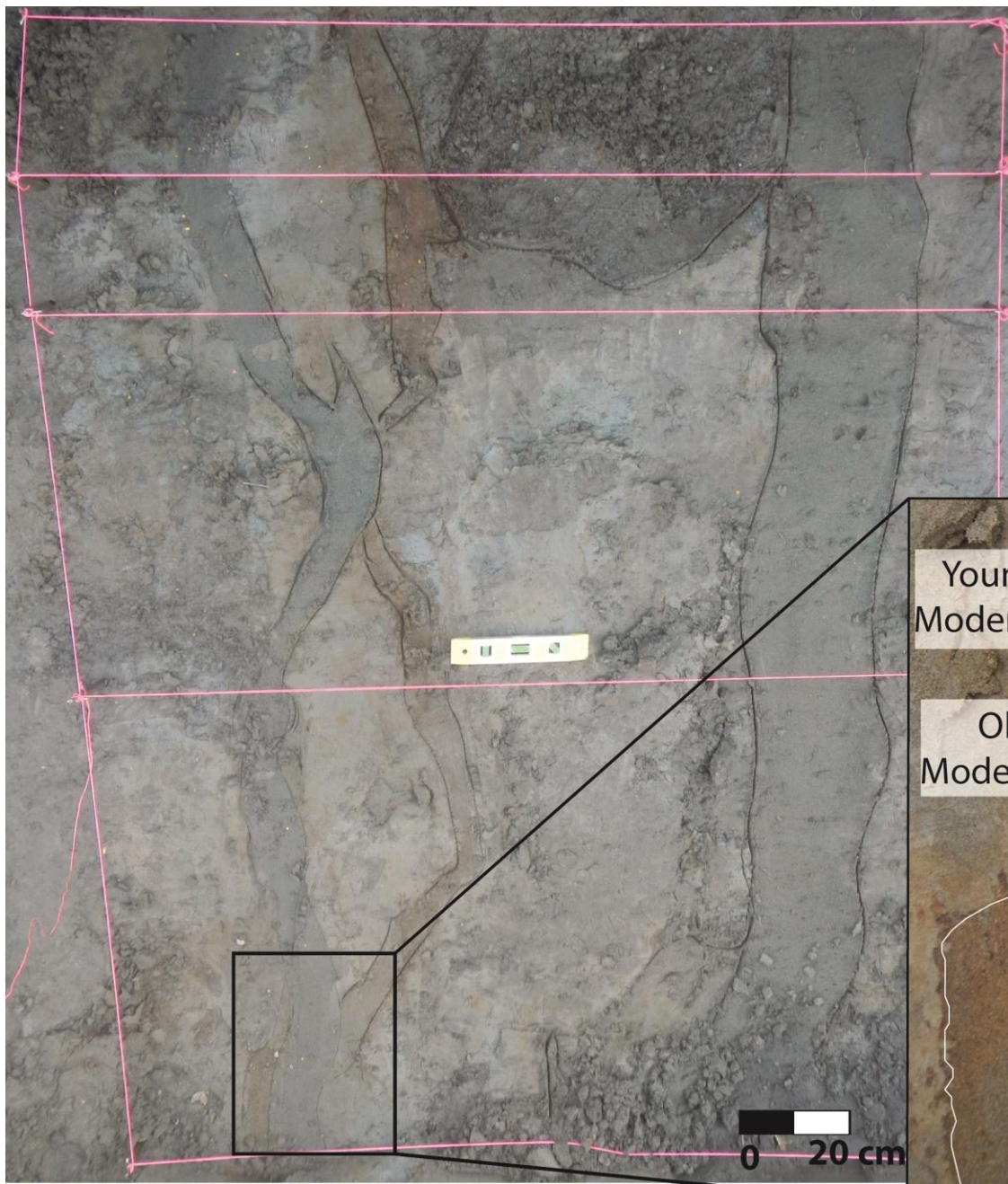
Crack length, orientation, dist from free face, elevation, etc

Targeted lateral spreading cracks for paleoliquefaction investigations

Difference between river and nearest downslope freeface azimuth



Geologic evidence for paleoliquefaction

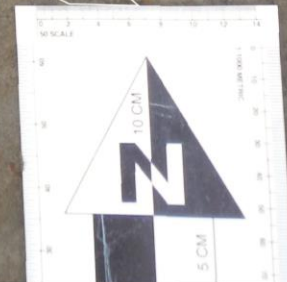


Youngest
Modern Dike

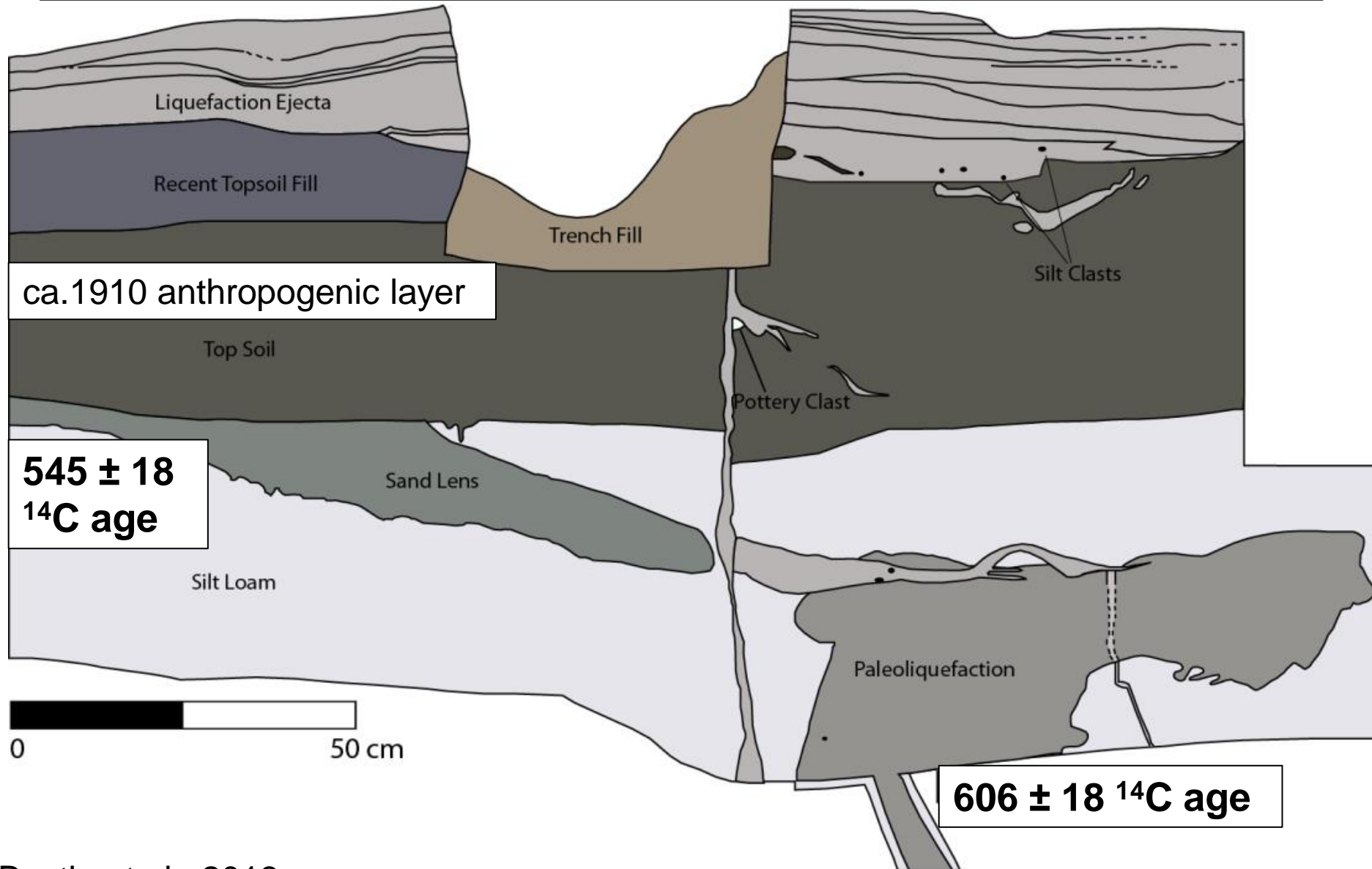
Older
Modern Dike

Paleo-dike

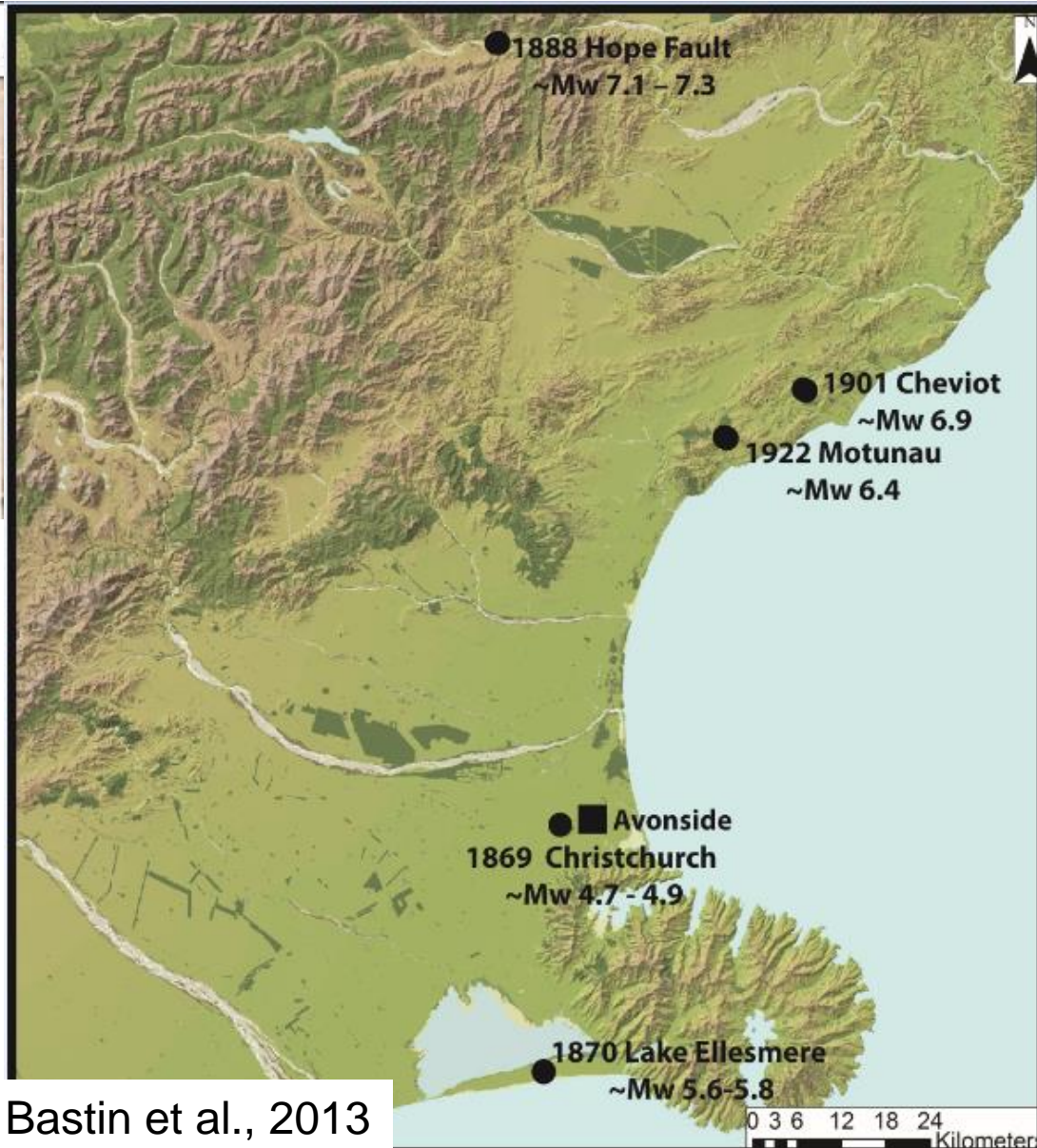
0 20 cm



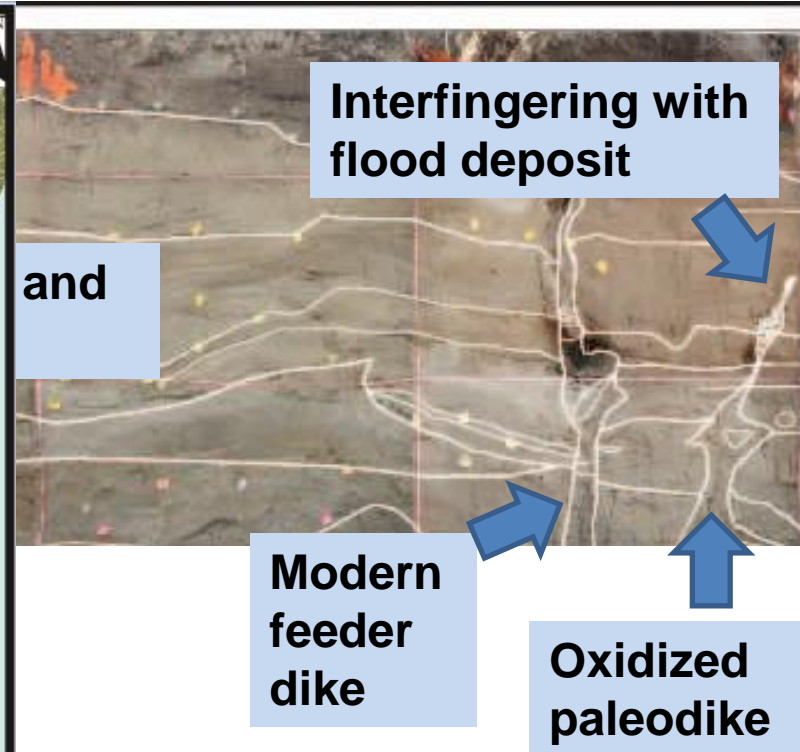
Liquefaction of my former backyard sometime between 1910 and ~1470 AD (last ~545 yr)



Liquefaction of Sullivan's Park sometime between 1880s and 1920s

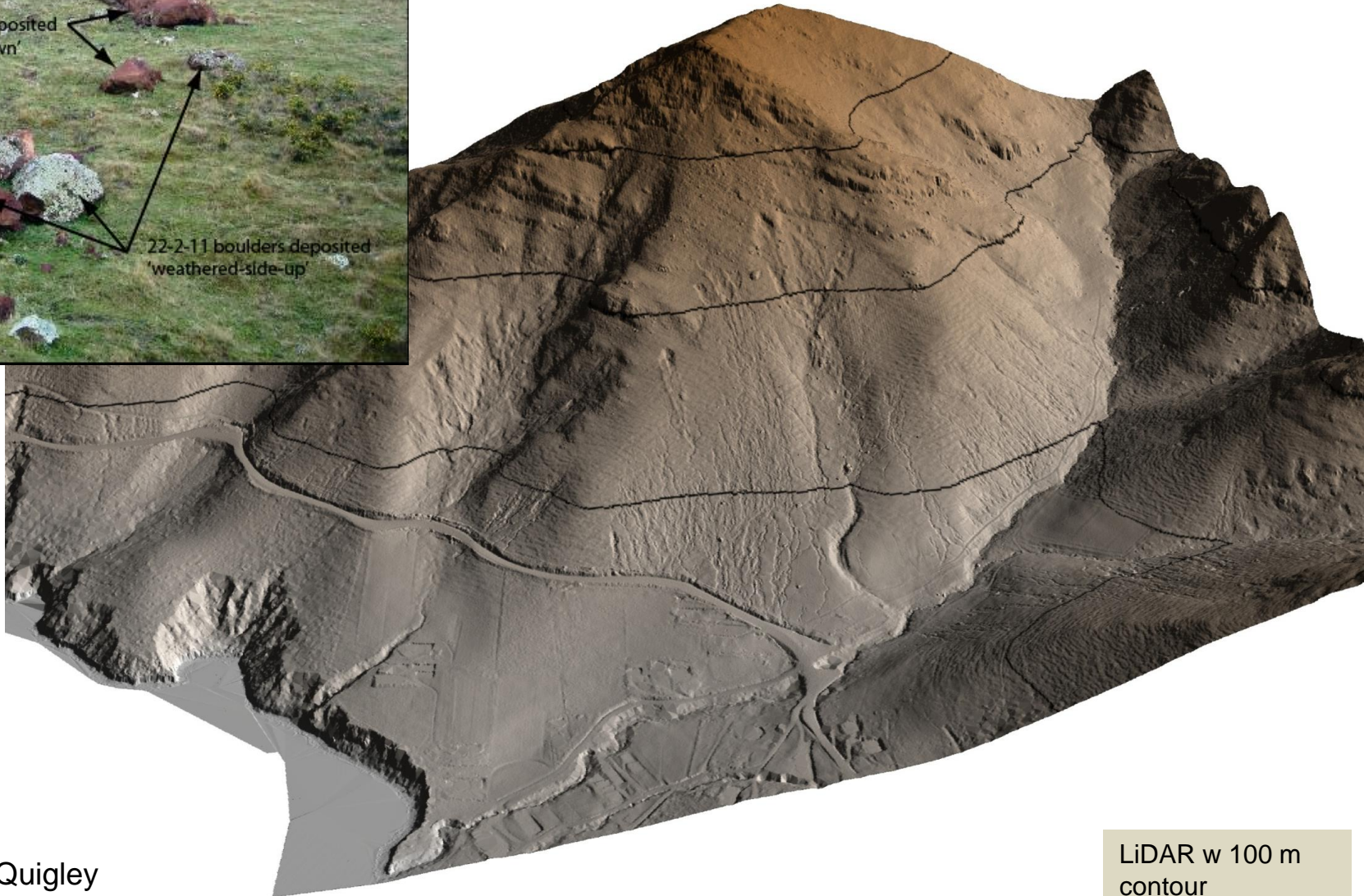


Bastin et al., 2013



Suspects

Rockfall / Boulder Roll Hazard

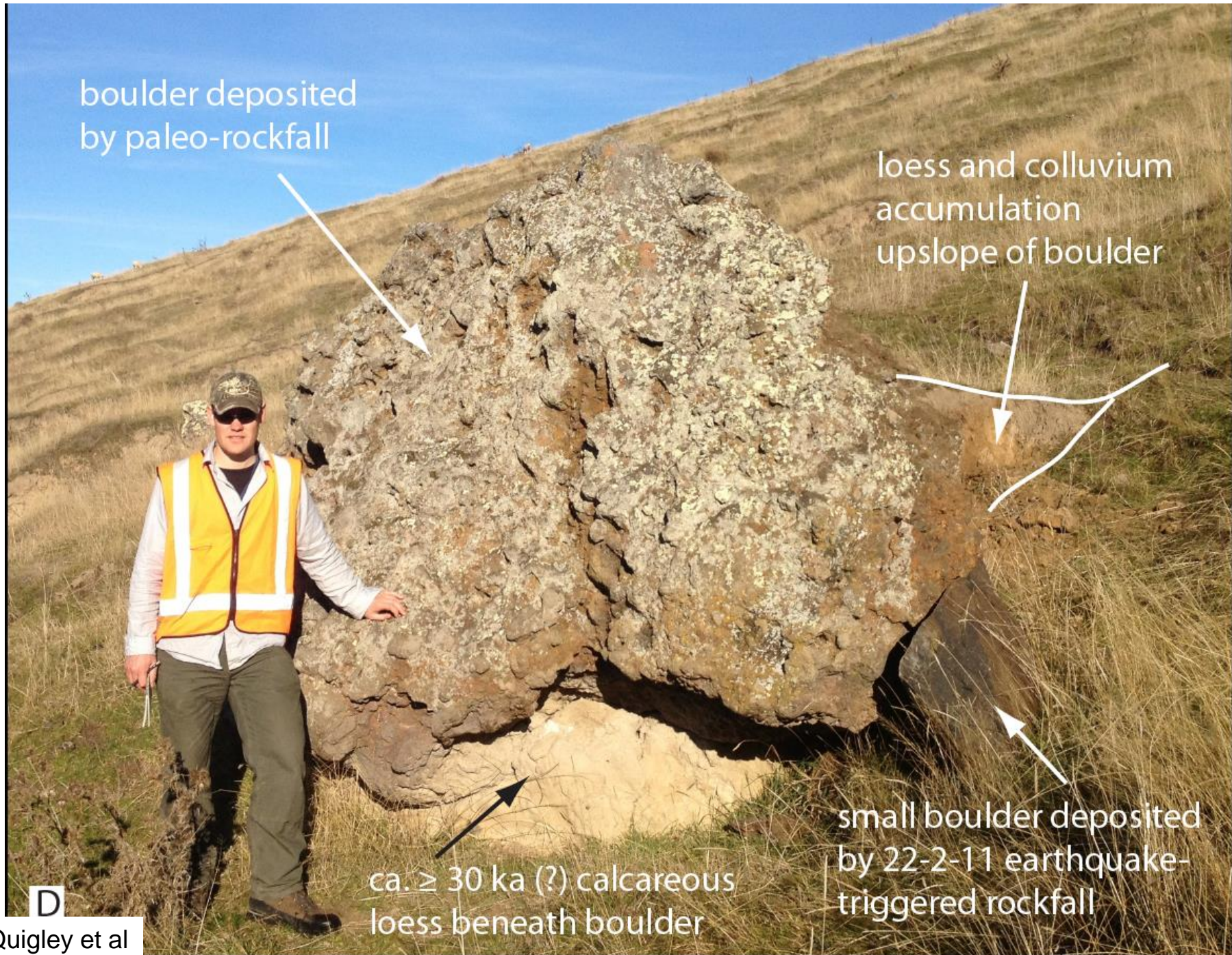


boulder deposited
by paleo-rockfall

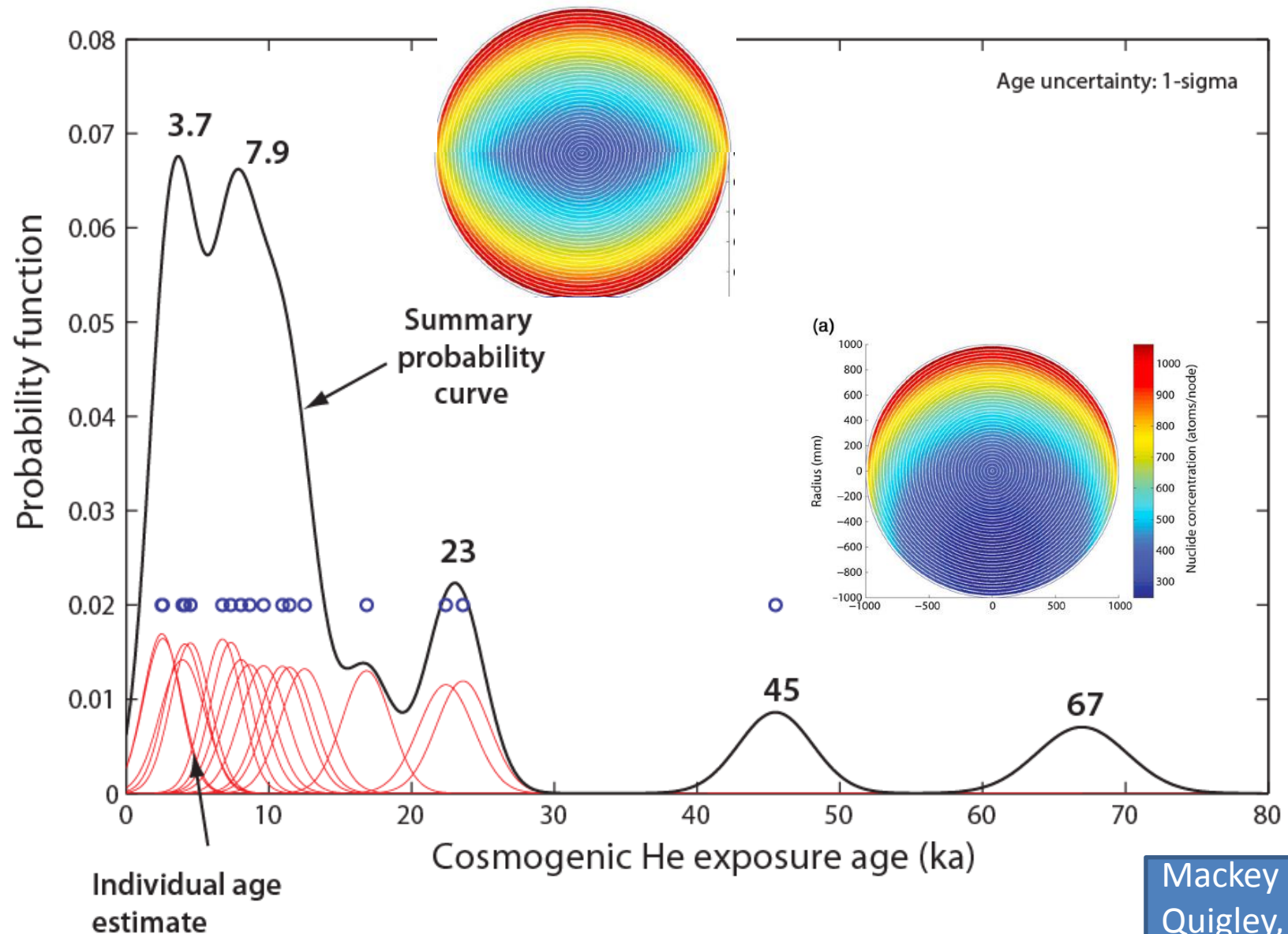
loess and colluvium
accumulation
upslope of boulder

ca. ≥ 30 ka (?) calcareous
loess beneath boulder

small boulder deposited
by 22-2-11 earthquake-
triggered rockfall



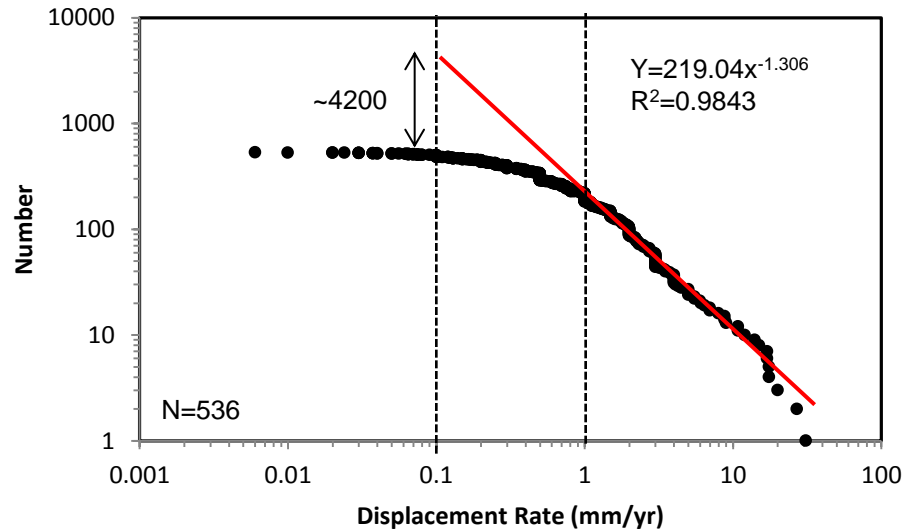
A mid Holocene rockfall event incorporating 'fresh' and pre-exposed material ?
No evidence for 'random' temporal occurrence
No evidence for Alpine Fault earthquakes



Future challenges

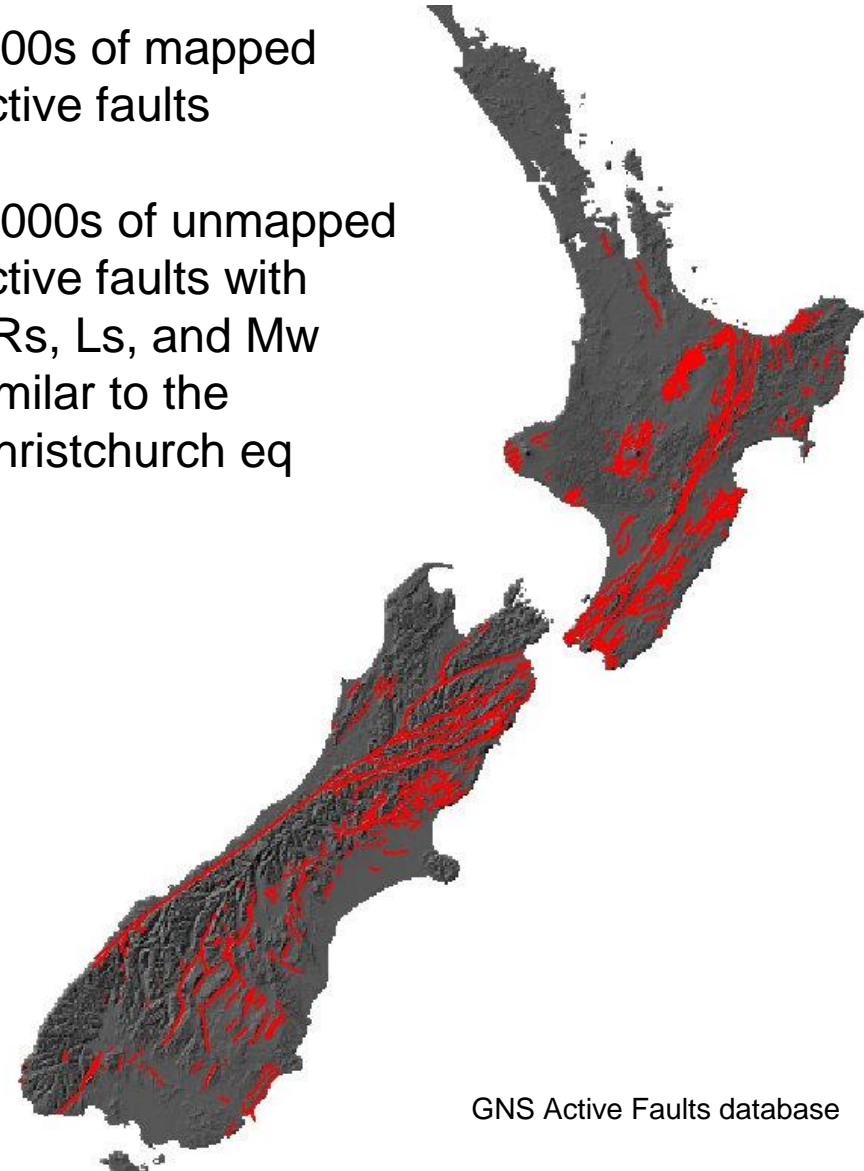
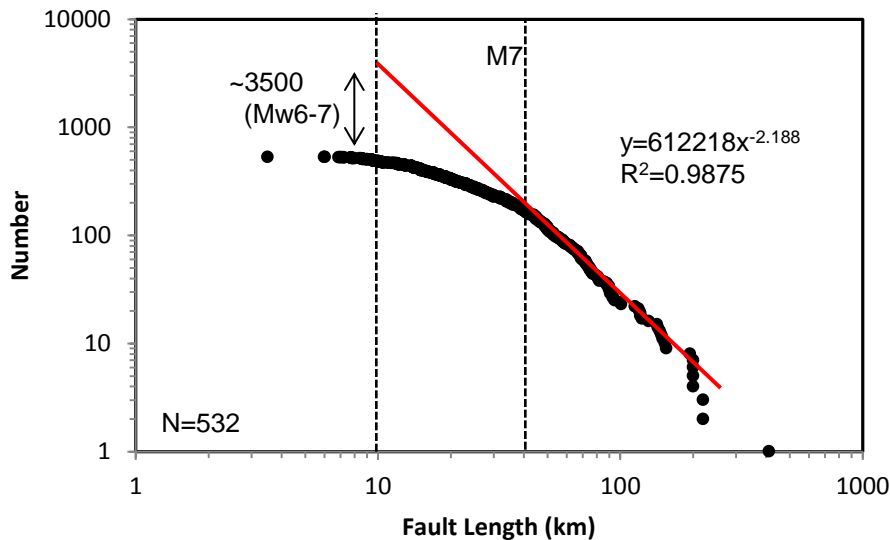


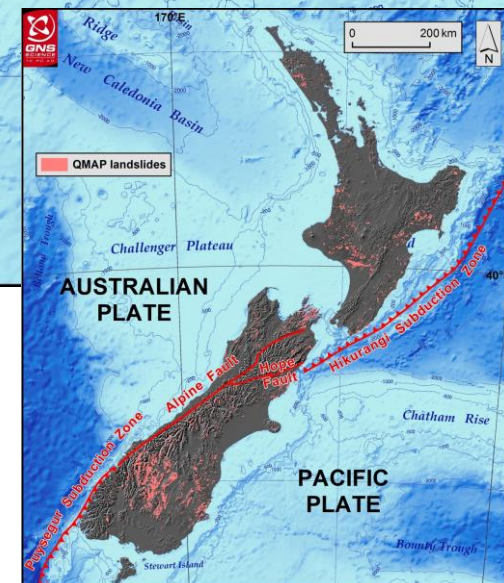
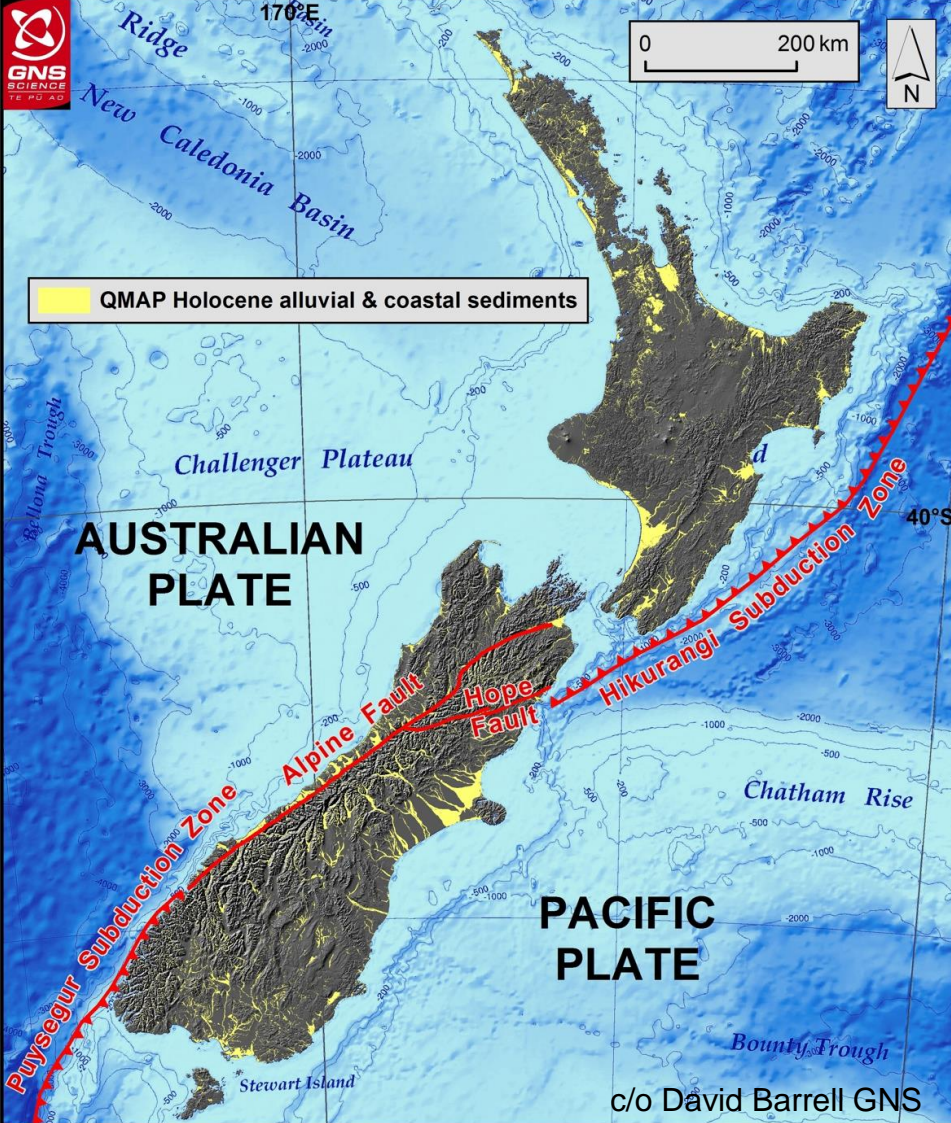
An analogous eq (sequence) could happen at any time anywhere else in New Zealand: have you personalized your hazard?



-100s of mapped active faults

-1000s of unmapped active faults with SRs, Ls, and Mw similar to the Christchurch eq





A land of potentially liquefiable sediments
and landslide-susceptible hillslopes

Conclusions

- Our research is driven by the desire to solve fundamental process-based questions and conduct ‘science for society’
- Earthquakes rupture the surface more than we recognize in the geologic record (*SRL* and *D*) – but discrete displacements provide meaningful M_w estimates when used with existing scaling relationships
- We see geologic evidence for penultimate earthquake rupture on the Greendale Fault ca. 25 ± 3 kyr ago, we infer from stress modelling that the complexity of this earthquake is the norm rather than the exception, we think it is unlikely that GF rupture could have propagated coseismically on to the Feb and June faults, but the fractures between these faults are effective ‘waveguides’
- Everywhere we look, we find geological evidence for pre-CES paleoliquefaction and paleorockfalls – are we listening to the geologic record?
- Personalize your hazard, and support proactive rather than reactive approaches to natural hazards